

# MERIS US Workshop

## 14 July 2008

### MERIS Level 2 processing

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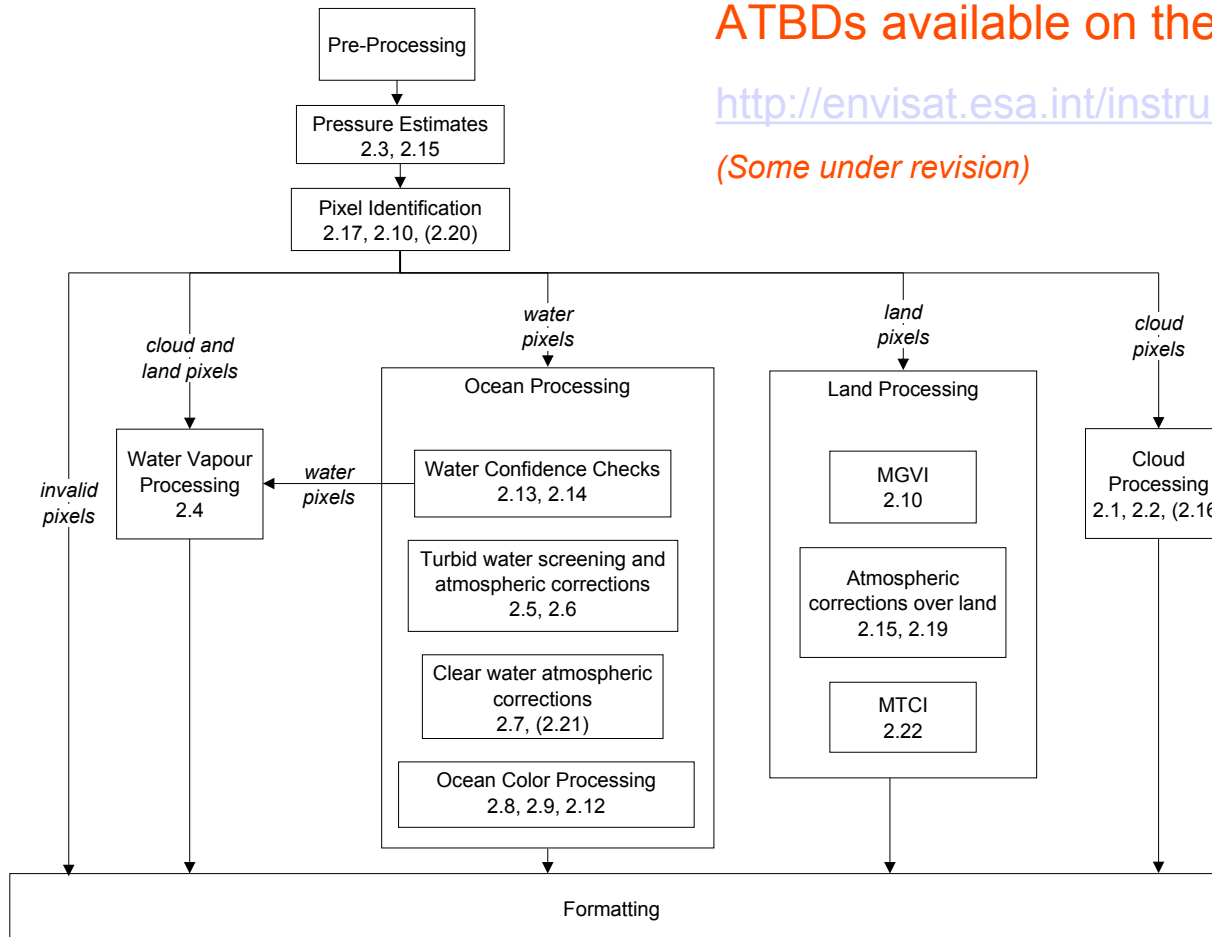
## Presentation Outlook

- PART 1: MERIS L2 PROCESSING GENERAL OVERVIEW
  - Common Processing
  - Cloud Branch
  - Land Branch
  - Water Branch
- PART 2: WATER BRANCH ALGORITHMS,  
A MORE DETAILED DESCRIPTION
  - Atmosphere correction
  - Ocean Color
- PART 3: TOWARD VICARIOUS ADJUSTMENT  
OF MERIS TOA REFLECTANCE (at Level 2)

# MERIS LEVEL 2 PROCESSING PART 1

## MERIS L2 PROCESSING GENERAL OVERVIEW

## Functionnal breakdown



ATBDs available on the web at:

<http://envisat.esa.int/instruments/meris/atbd/>

*(Some under revision)*

MERIS ESL:

Institutes:

- FUB
- GKSS
- JRC
- LISE/ULCO
- LOV
- PML

Companies:

- ACRI
- BC

## Pre-processing

- For **all** pixels:
  - Interpolate all annotations at pixel
  - Derive *surface* pressure from ECMWF *mean sea level* and pixel altitude
- For **valid** pixels only:
  - convert L1b radiance  $L$  into TOA reflectance  $\rho$

## Pressure processing

From

- measurement within the O<sub>2</sub>-A absorption line at 761 nm (band 11) and reference at 753 nm (band 10)
- accurate knowledge of the band 11 central wavelength
  - ➔ Cloud Top Pressure (neural net)
  - ➔ Surface Pressure (polynomials)
- Derived at every valid pixel (input to classification)
- Valid product only over relevant surface types (cloud and land pixels, respectively)

## Pixel identification

- Goal: sort valid pixels into Cloud, Water and Land
- Means:
  - Cloud screening:
    - spectrum values & slopes (on Rayleigh corrected reflectance)
    - pressure tests
  - Land/water discrimination:
    - *A priori* land/water classification
    - radiometric (clear sky) land/water reclassification where a *priori* is questionable (on gas corrected reflectance)

## Smile Correction

- “Smile”: in-FOV variation of channels central wavelength
- Can affect Level 2 products if not accounted for
- Need for a simple and robust smile correction
  - First order correction  $\rho_{ng}^* = \rho_{ng} + \frac{\partial \rho}{\partial \lambda} \cdot \Delta \lambda$
  - Restricted to Land and Water pixels
  - Restricted to surface dependent subset of bands
  - Based on surface dependent spectral slope estimates



## Cloud branch

- Compute the ***cloud albedo, cloud optical thickness, cloud type*** products (in addition to already available ***cloud top pressure***)
- Albedo and optical thickness are derived from polynomials of  $L_{TOA}(753)$  using geometry and surface albedo dependant coefficients. Polynomial coefficients are derived from radiative transfer simulations.
- *Cloud type* product is an *index* computed from CTP and CA (defined by ISCCP climatology project)

## Total Column Water Vapour

- Determine column water vapour content above all surfaces from absorption at 900 nm
- Polynomials of  $T = L_{\text{TOA}}(900) / L_{\text{TOA}}(885)$  (from RT simulations)
- Surface dependent algorithms
  - Above land and water with high glint:
    - Correct for surface reflectance and its spectral dependency, estimated from  $L_T$  at 753, 885, 900
    - Account for surface pressure (vertical profile)
  - Above water, no/low glint :
    - Correct T for aerosol estimated from  $L_T$  at 779, 865
  - Above cloud :
    - Account for cloud optical thickness (a MERIS cloud product)
    - Account for underlying surface albedo

# Land branch

MERIS  
Global Vegetation Index

Atmosphere Corrections  
over Land



MERIS  
Terrestrial Chlorophyll Index

## Land branch

MERIS Global Vegetation Index



- Atmosphere & geometry insensitive vegetation index
- Remove unwanted pixels → **flags**
- Based on **normalised**  $\rho_{TOA}(B,R,IR)$  (→ anisotropic)
- Computes **rectified**  $\rho_R$  for R & IR (→ top of canopy at ref geometry)
- Computes **MGVI**( $\rho_R(R), \rho_R(IR)$ ) (→  $\equiv$ FAPAR)

Atmosphere Corrections over Land



MERIS Terrestrial Chlorophyll Index

## Land branch

MERIS Global Vegetation Index

Atmosphere Corrections over Land

MERIS Terrestrial Chlorophyll Index

- compute top of aerosol  $\rho_{top}$  (all pixels)
- Retrieve aerosol over vegetation
  - select DDV pixels using ARVI
  - biome climatology & models  
→  $\rho_S$  at 3 bands
  - find best aerosol model  $\alpha$  and  $\tau_{443}$  so that  $\{\rho_S\}_{(3 \text{ bands})}$  propagated to TOA match measures

## Land branch

MERIS Global Vegetation Index

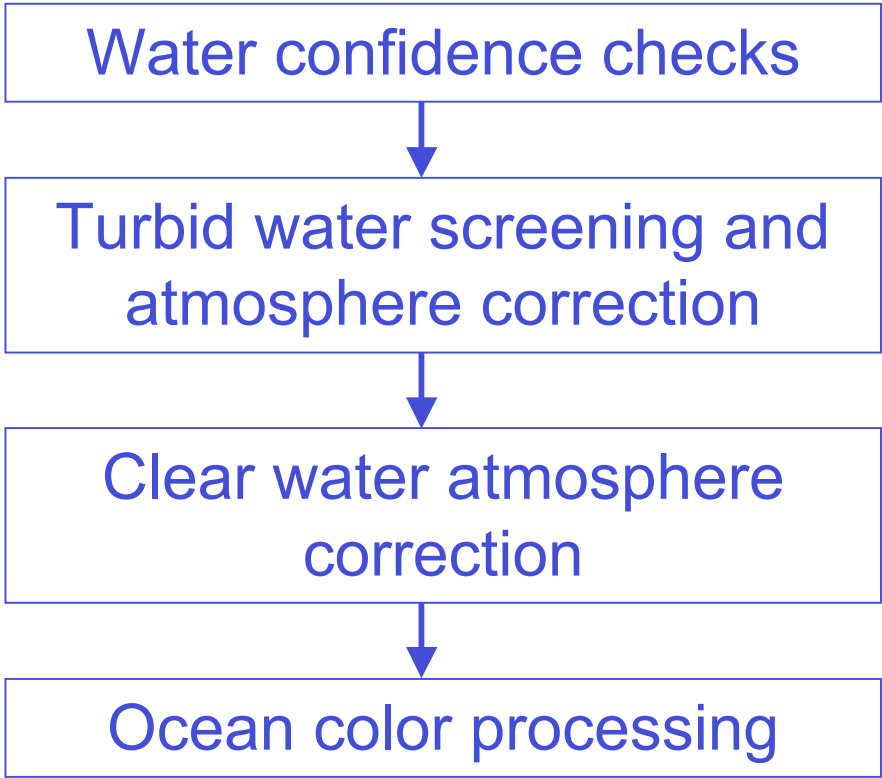
Atmosphere Corrections over Land

MERIS Terrestrial Chlorophyll Index

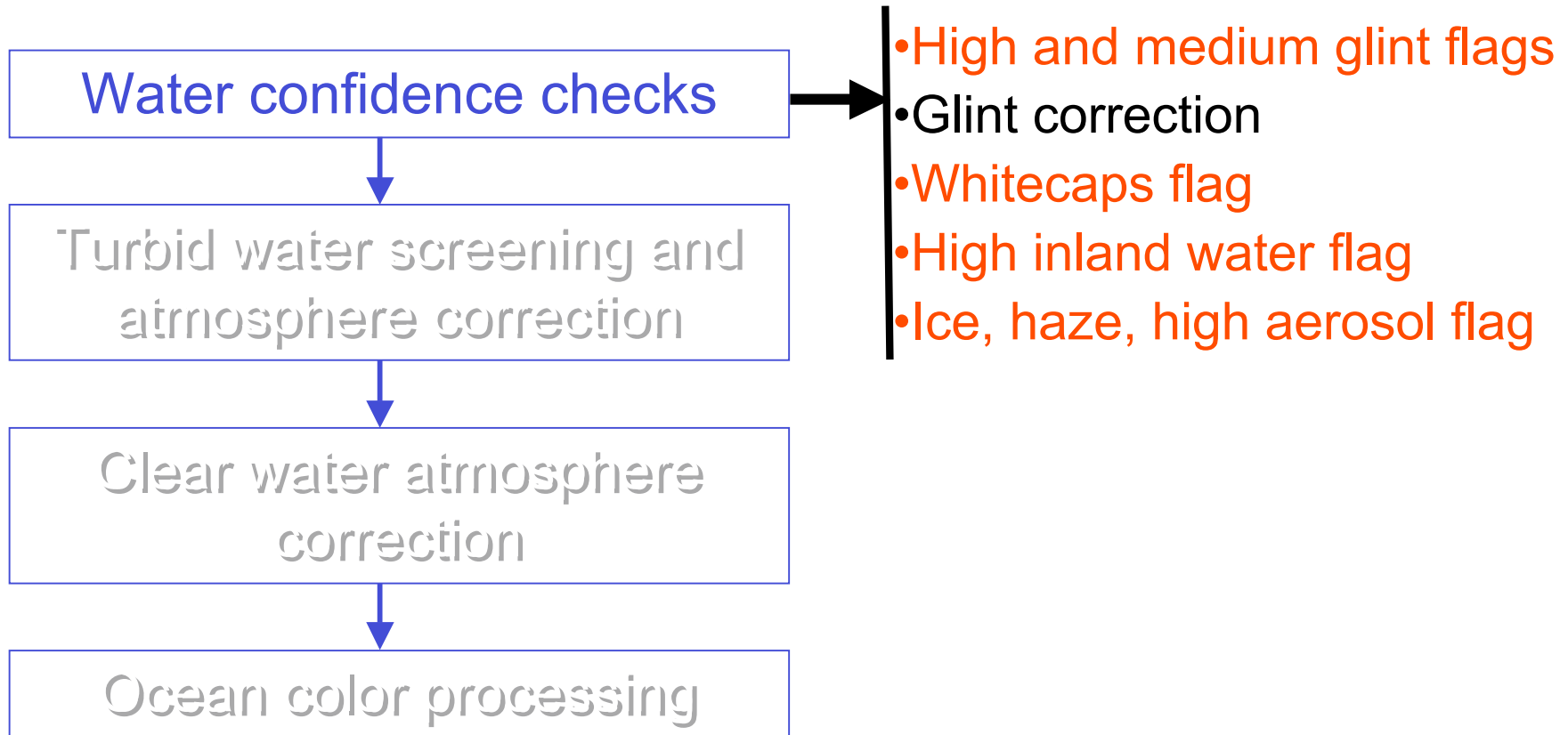
- Remove unwanted pixels → flags
- Based on  $\rho_{TOP}(R, IR1, IR2)$

$$\bullet \text{ MTCI} = \frac{\rho_{TOP}(753) - \rho_{TOP}(709)}{\rho_{TOP}(709) - \rho_{TOP}(681)}$$

# Water branch



## Water branch





## Water branch

Water confidence checks



Turbid water screening and atmosphere correction



- Identify bright waters (sediment dominated case 2, **CASE2\_S flag**)
- Estimate sediment load
- Estimate IR marine signal  $\rho_w$

Clear water atmosphere correction



Ocean color processing

## Water branch

Water confidence checks



Turbid water screening and atmosphere correction



Clear water atmosphere correction



- Identify aerosol:  $\tau_{865}$ ,  $\alpha_{NIR}$  (from IR where  $\rho_w$  known)
- Estimate atmosphere path reflectance including molecular/aerosol coupling
- Provide  $\rho_w$  (all b but 761 & 900)
- **Flags**: quality and science



Ocean color processing

## Water branch

Water confidence checks



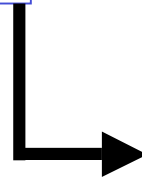
Turbid water screening and atmosphere correction



Clear water atmosphere correction



Ocean color processing



From  $\rho_w$ , provide:

- Algal 1 (case 1)
- Algal 2, TSM, YS (case 2)

• Instantaneous Photosynthetically Available Radiation (PAR)

# MERIS LEVEL 2 PROCESSING PART 2

## A MORE DETAILED DESCRIPTION OF WATER BRANCH ALGORITHMS

## Water Processing: more details

- Input: top of atmosphere reflectance
- Mains steps:
  - Atmosphere correction
    - Gas correction ( $O_3$ ,  $H_2O$ ,  $O_2$ )
    - Pixels screening & glint
    - Coupled molecules/aerosols correction
    - directional water leaving reflectance
  - Ocean color processing
    - Chl 1 from band ratio (case 1)
    - Chl 2, TSM, YS from Neural network (case 2)

$$\rho_w = \pi \frac{L_w}{E_d^{(0+)}}$$

## Water processing: atmosphere model

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[stratospheric aerosol layer]

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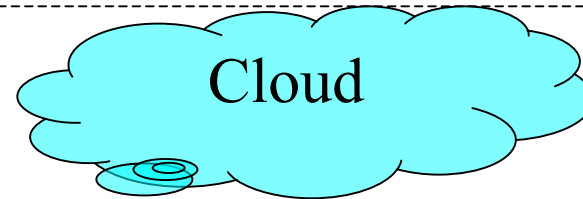
ozone layer (absorption)

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O<sub>2</sub>, H<sub>2</sub>O (absorption coupled with scattering)

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Air + aerosol  
mixture

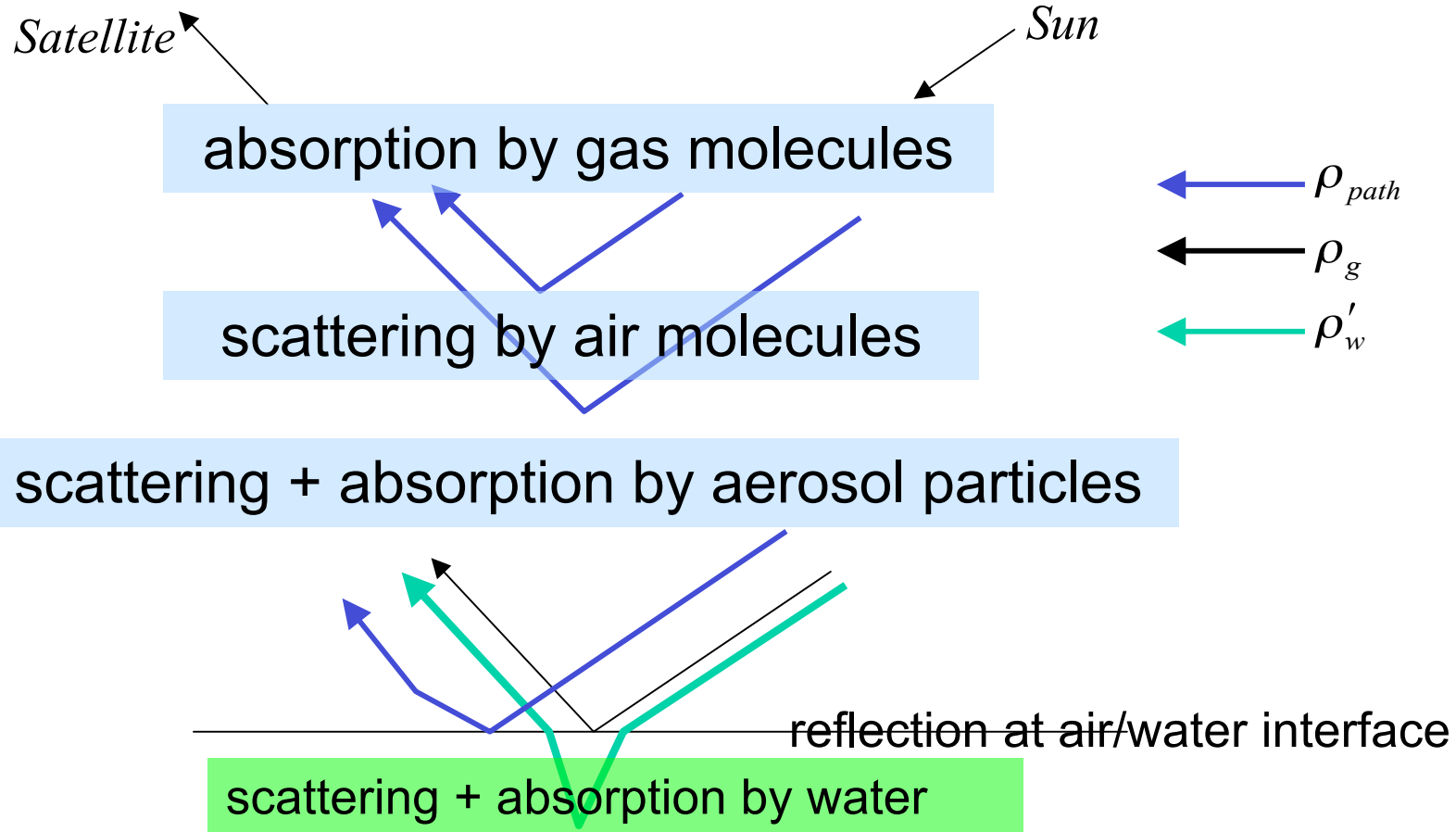


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Air-water interface

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# Water processing: simplified signal model



$$\rho = t_g \cdot (t_{dir} \cdot \rho_G + \rho_{path} + t_{up} \cdot T_d \cdot \rho_w)$$

# Water Processing: gas correction

## Gas corrections

- O<sub>3</sub>: transmittance  $t_{O_3}$  from ECMWF ozone and optical thickness for each band
- H<sub>2</sub>O: transmittance at any band  $t_{H_2O}(b)$  estimated from measured  $t_{H_2O}(900)$
- O<sub>2</sub>: transmittance at any band  $t_{O_2}(b)$  estimated from measured  $t_{O_2}(761)$



## Water Processing: glint & bright pixels

### Pixels screening & glint flagging/correction:

- Glint reflectance estimates: Cox & Munk anisotropic model (revised with Ebushi & al parameters, and fed by ECMWF wind vectors)
- 2 glint thresholds and flags:
  - high glint: too high for reliable correction ( $\rho_G \geq \alpha \cdot \rho(865)$ )
  - medium glint: correction is required and possible ( $\epsilon \leq \rho_G < \alpha \cdot \rho(865)$ )  $\rightarrow \rho_{GC}(\lambda) = \rho_{ng}(\lambda) - t \cdot \rho_G$
- “Bright pixels” screening: undetected clouds, haze, high aerosols and ice flagged (from a reflectance test at 412)
- Whitecaps: raise flag upon wind speed, no correction

## Water Processing: path reflectance

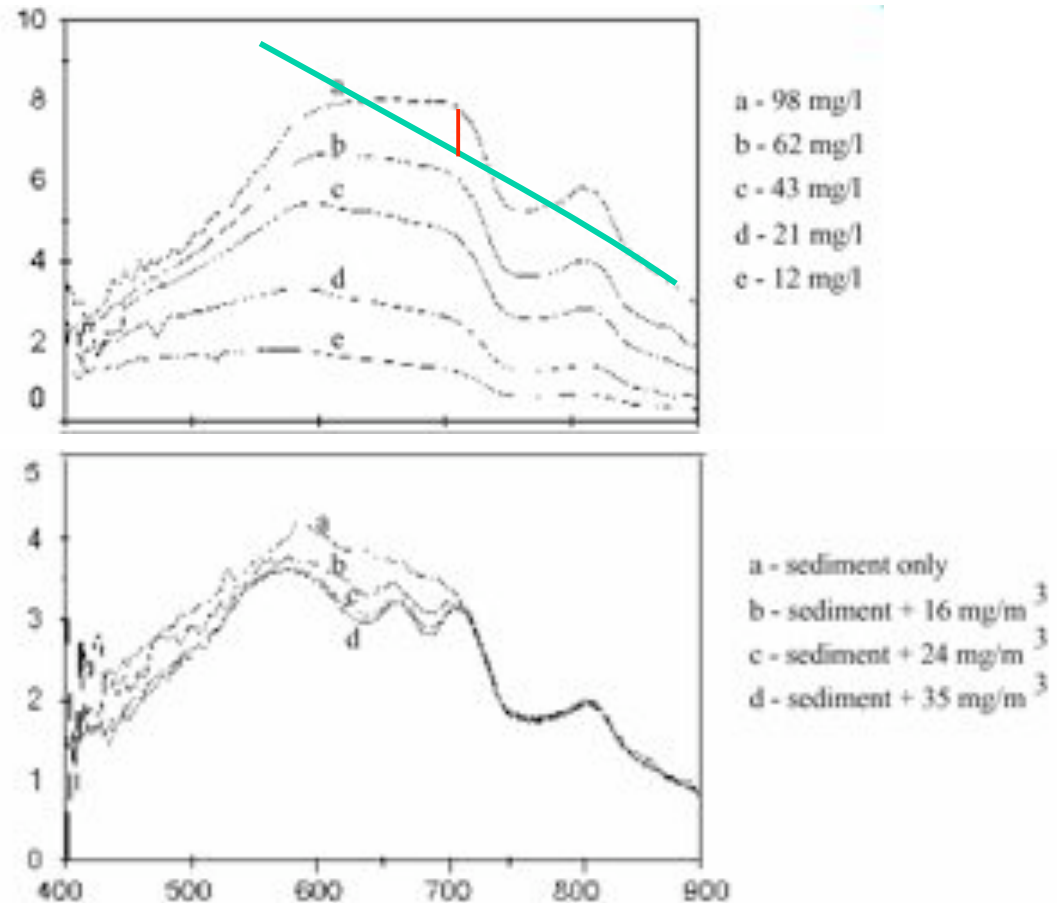
- Based on a classical “black ocean in the NIR” approach
- Using standard aerosol models closed to SeaWiFS ones
- With some specificities:
  - Rayleigh/aerosol coupling accounted for through the use of  $\rho_{\text{path}}/\rho_{\text{R}}$  (ATBD 2.7)
  - Extension to (moderately) sediment loaded waters through estimation of  $t.\rho_{\text{w}}$  in the NIR (ATBD 2.6)
  - Additional aerosols with steep spectral dependency
  - Detection of absorbing aerosols using “deficit of reflectance at 560 nm” (ATBD 2.7) and corresponding additional models (Moulin et al)

# Water Processing: turbid waters

Purpose:

To identify waters with significant backscatter in the NIR

To provide corresponding reflectance estimates allowing atmosphere corrections



# Water Processing: turbid waters

- Simplified atmospheric model :  
single scattering approximation

$$\rho_{rc} = \rho_{GC} - \rho_R \text{ so that } \rho_{rc} - t\rho_w = \rho_a$$

$t$  = diffuse transmission up & down, Rayleigh only

$$\rho_{GC} = \rho_R + \rho_a + t \cdot \rho_w$$

$$\frac{\rho_a(b1)}{\rho_a(b2)} = \left( \frac{\lambda(b1)}{\lambda(b2)} \right)^n$$

- Find  $n$  and TSM  
that explain  $\rho_{RC}$   
at 709, 779 & 865

$$\left\{ \begin{array}{l} \rho_{rc}(865) - t\rho_w(865) = [\rho_{rc}(779) - t\rho_w(779)] \cdot \left( \frac{\lambda(779)}{\lambda(865)} \right)^n \\ \rho_{rc}(865) - t\rho_w(865) = [\rho_{rc}(709) - t\rho_w(709)] \cdot \left( \frac{\lambda(709)}{\lambda(865)} \right)^n \end{array} \right.$$

- Realized using a simplified  
 $\rho_w$  model

$$\rho_w = \pi \cdot \mathfrak{K} \cdot f/Q \cdot \frac{b_{bw} + b_{bp}^* \cdot TSM}{a_w}$$

- Ends up with  $t \cdot \rho_w$  estimates at 779, 865 + CASE2\_S flag  
(ATBD 2.6 under revision)

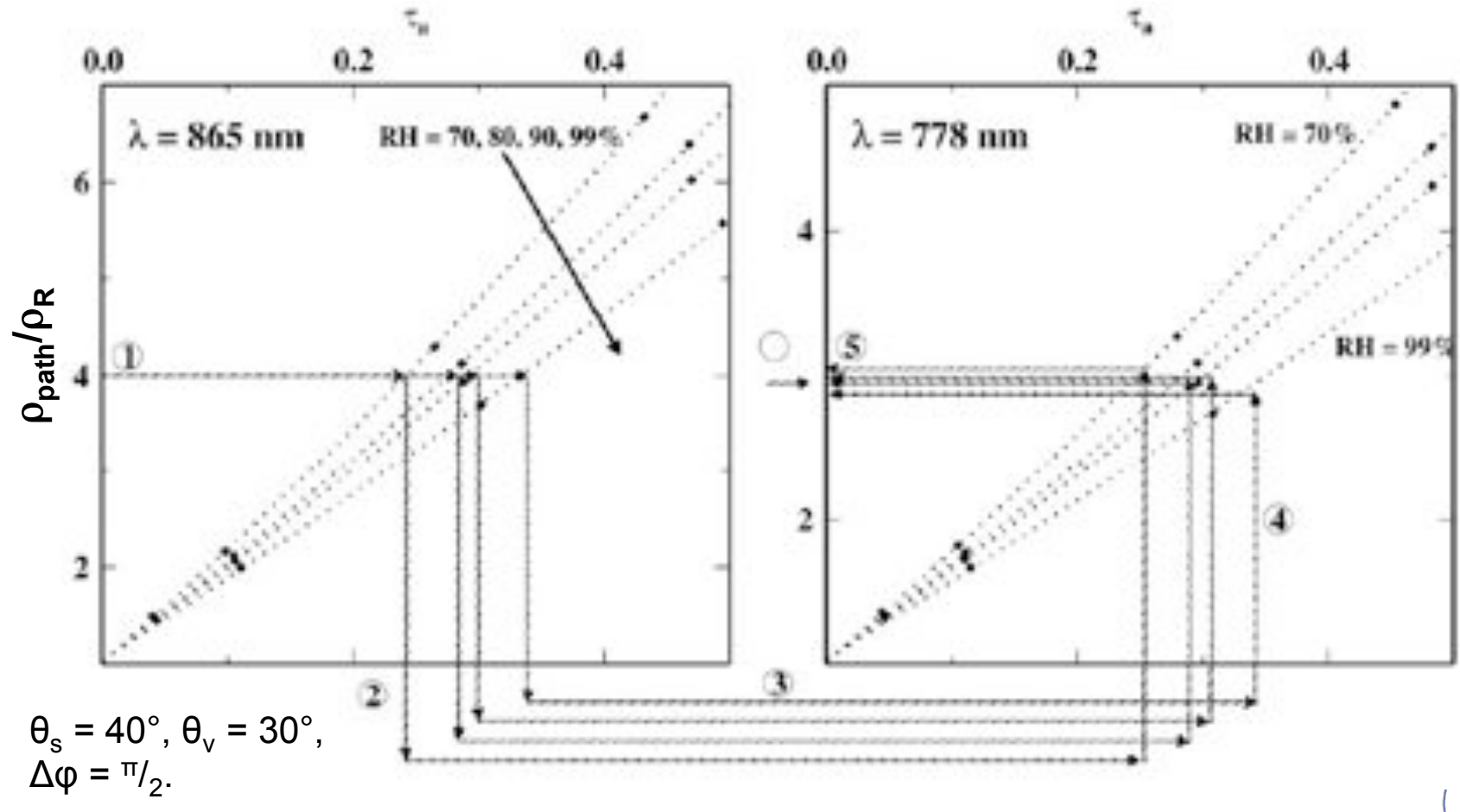
# Water Processing: aerosol identification

## Principle:

- Inputs:  $\rho_{\text{path}} = \rho_{\text{GC}} - t \cdot \rho_w$  at 779 and 865 ( $= \rho_{\text{GC}}$  in case 1)
- Tabulated relationships:
  - $\rho_{\text{path}}/\rho_R = f(\tau, \lambda, \Omega, \text{model}), \quad (1)$
  - $\tau = f^{-1}(\rho_{\text{path}}/\rho_R, \lambda, \Omega, \text{model}) \quad (2)$
  - $\tau(\lambda, \text{model}) = \tau(\lambda_{\text{ref}}, \text{model}) \quad (3)$
- For all aerosol models:
  - Measured  $\rho_{\text{path}}/\rho_R$  at 865 gives  $\tau_{865}$  using (2)
  - $\tau_{865}$  gives  $\tau_{779}$  using (3)
  - $\tau_{779}$  gives  $\rho_{\text{path}}/\rho_R$  at 779 using (1)
- Select pair of models most closely bracketing measured  $\rho_{\text{path}}/\rho_R$  at 779 and mixing ratio that fits.

# Water Processing: aerosol identification

Illustration with a set of 4 maritime models



# Water Processing: aerosol identification

## Global procedure:

- identification done over the whole set of non-absorbing aerosols (MAR, COA, RUR with various RH + 3 “blue”)

- best pair is selected, derive:  $\rho_{path}^{calc} = \left( \frac{\rho_{path}}{\rho_R} \right)^{calc} \cdot \rho_R$  at 510 nm

- If not CASE2\_S then:

- Get  $\rho_w(510)$  using  $\rho_w = \frac{\rho_{GC} - \rho_{path}^{calc}}{t_{up} \cdot t_{down}}$

- If  $\rho_w < (\text{mean} + \sigma)_{climatology}$  repeat procedure with the absorbing aerosol families  $\rightarrow$  new candidate pairs

- Select the one that minimizes  $|\rho_w - (\text{mean} + \sigma)_{climatology}|$

- Proceed to correction at all bands

## Water Processing: aerosol products

- Aerosol optical thickness at 865nm
- Aerosol Angstrom exponent (779-865)
- Atmosphere correction flags:
  - validity flag for aerosol products
  - OADB (out of aerosol database: limited representativity)
  - ABSO\_D: absorbing aerosols were selected
  - CASE2\_S: significant scattering by water in the NIR
  - Medium glint
  - High glint
  - Ice/haze/high aerosol



## Water Processing: Ocean Color

### Ocean color products:

- Algal 1 index ( $\sim\text{Chl}_a$  concentration,  $\text{mg}\cdot\text{m}^{-3}$ ): case 1
- Algal 2 index ( $\sim\text{Chl}_a$  concentration,  $\text{mg}\cdot\text{m}^{-3}$ ): case 2
- Total Suspended matter ( $\text{g}\cdot\text{m}^{-3}$ ): case 2
- Yellow Substance (absorption at 442,  $\text{m}^{-1}$ ): case 2
- Flags:
  - Validity flags for each product
  - CASE2\_Anomalous: excess of reflectance at 560 as compared to derived Algal1  $\rightarrow$  anomalous scattering

## Water Processing: Algal 1

- In Case 1 water,
  - $R(\lambda)$  is a function of algal pigments
  - $R(442)/R(560)$  (or  $R(490)/R(560)$ ) is directly related with Chl

$$R = \frac{E_u}{E_d}$$

$E_u$  : upwelling irradiance (integral of radiance) just below the surface

$E_d$  : downwelling irradiance

$$\rho_w = \pi \cdot \mathfrak{R} \cdot \frac{R}{Q}$$

contribution of the air-water interface

bi-directional distribution of the upwelling radiance  
(*depends on Chl*)

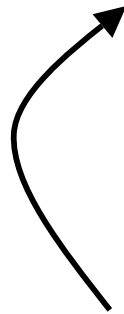
From MERIS case 1 model: 
$$\rho_w = \pi \cdot \mathfrak{R}(\theta_s, \theta_v, \tau, w) \cdot \frac{f(\lambda, \theta_s, \tau, w, Chl)}{Q(\lambda, \Omega, \tau, w, Chl)} \cdot \left[ \frac{b_b}{a} \right] (\lambda, Chl)$$

## Water Processing: Algal 1

- Reflectance correction for bi-directionality requires knowledge of Chl → iterative method:

- Select highest  $\frac{\rho'_w(b_i)}{\rho'_w(b_{ref})}$  ratio among  $b_i \in \{443, 490, 520\}$ ,  $b_{ref} = 560$
- Normalization of  $\rho_w$  ratio using tabulated  $f/Q$  (Chl dependent)

Iterative process



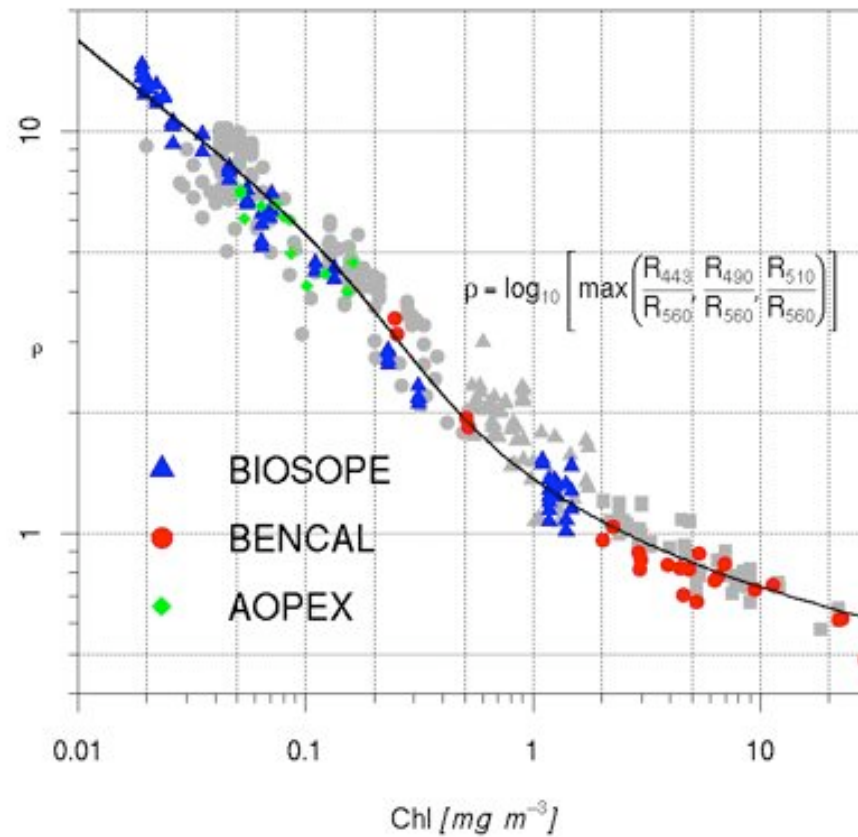
$$band\_ratio = \frac{\rho'_w(b_i)}{\rho'_w(b_{ref})} \cdot \frac{Q_i}{Q_{ref}} \left( = \frac{f/Q(b_{ref})}{f/Q(b_i)} \cdot \frac{\rho'_w(b_i)}{\rho'_w(b_{ref})} \cdot \frac{f_0(b_i)}{f_0(b_{ref})} \right)$$

- $\log \text{Chl1} = \text{polynomial}(\text{band\_ratio})$

# Water Processing: Algal 1

OC4Me

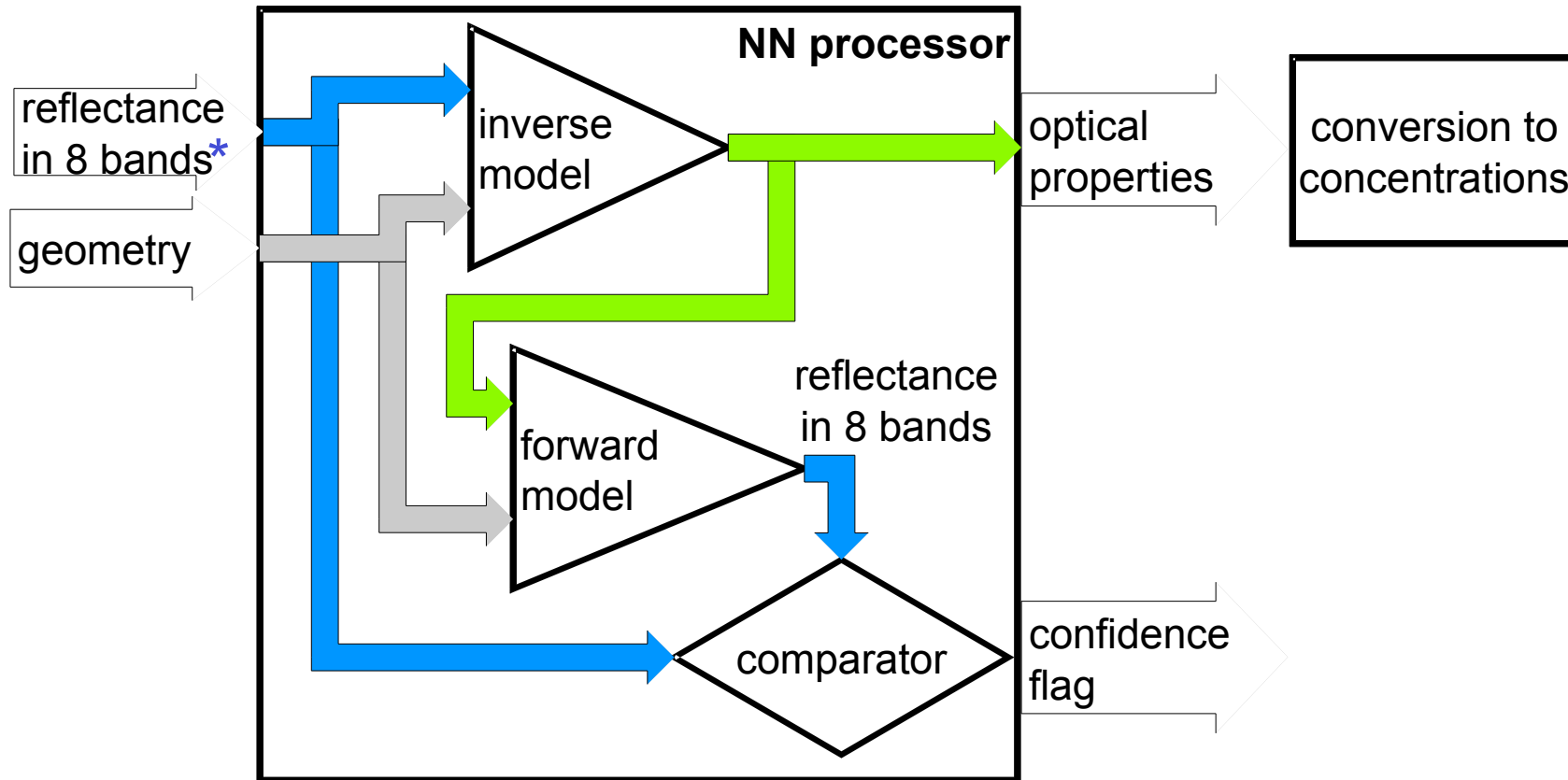
$$\text{Chl} = 10^{(0.40657) + (-3.6303) \rho + (5.44357) \rho^2 + (-5.48061) \rho^3 + (1.75312) \rho^4}$$



## Water Processing: Case 2 Neural Network

- Case 2 waters:
  - Described by *optical components*:
    - Phytoplankton pigment absorption  $a_{\text{pig}}$
    - Absorption of all other substances  $a_{\text{yp}}$   
(dissolved Gelbstoff  $y$  and non-algal particles  $p$ )
    - Scattering of all particles  $b_p$
  - Global inversion is done using a Neural Network (NN), provides  $a_{\text{pig}}(442)$ ,  $a_{\text{yp}}(442)$ ,  $b_p(442)$  from reflectance in 8 bands
  - NN is trained offline over a large data set produced by Radiative Transfer model
  - Internal consistency verification with a forward model NN
  - Components converted to concentrations using reversible relationships

## Water Processing: Case 2 Neural Network



\*: 412 to 709, excluding 681

## Water Processing: Case 2 Neural Network

- Neural Net principle and training conditions are described in a recent publication:
  - Doerffer, R. and Schiller H. (2007) 'The MERIS Case 2 water algorithm', Internal Journal of Remote Sensing, 28:3, 517-535
- As well as:
  - MERIS ATBD 2.12 (*under revision*)
  - MERIS Reference Model Document (*under revision*)

# MERIS LEVEL 2 PROCESSING PART 3

## TOWARD VICARIOUS ADJUSTMENT OF MERIS TOA REFLECTANCE



## Vicarious Adjustment: Principle

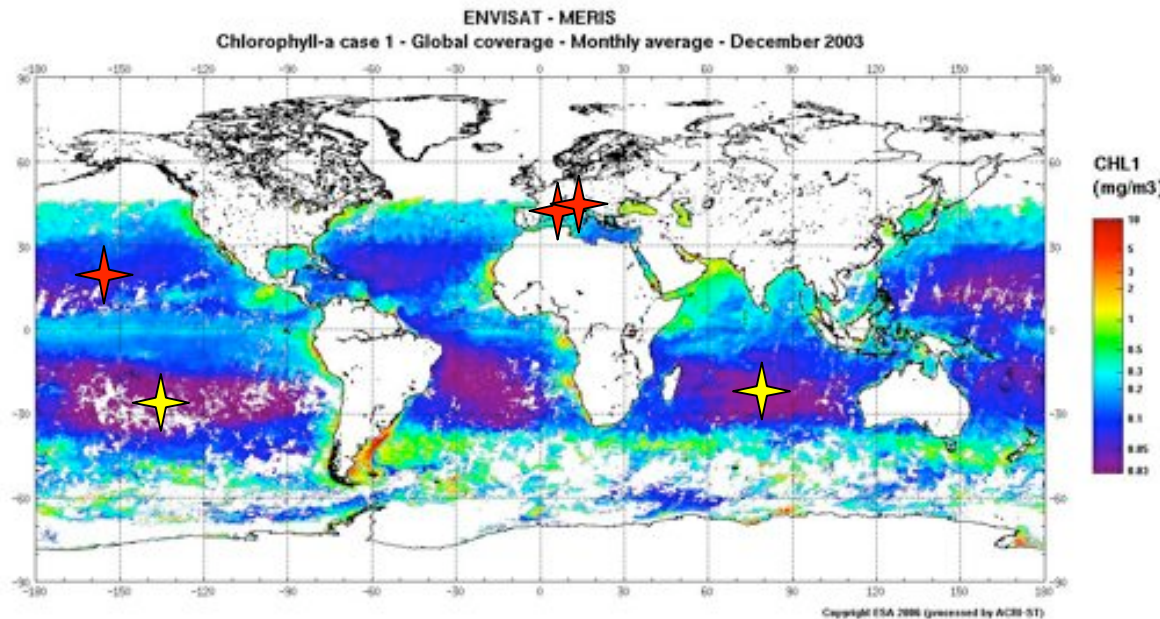
- Based on the work currently led at NASA for SeaWiFS and MODIS vicarious calibration, see Franz et al. 2007, Bailey et al. 2007.
- Consist in computing averaged multiplicative gain factors to correct the “TOA signal”, thanks to a DB of reference in-situ  $\rho_w^t$ 
  - “TOA signal” = Level 2 reflectance pre-corrected for smile, gaseous absorption, glint, i.e. at input of the Water AC algorithm
  - $\rho_{gc}^{new}(\lambda) = \rho_{gc}(\lambda) * G(\lambda)$  for  $\lambda$  in the VIS and NIR
- Two-step approach separating the NIR and VIS channels, avoiding iterative procedure within the AC algorithm
  1. First adjust one of the two NIR bands used in aerosol retrieval (the other one assumed perfect)  $\rightarrow G(779)$  for MERIS AC
  2. AC being now assumed “perfect”, adjust all other bands  $\rightarrow G(\lambda)$
- The methodology fully imbricates the sensor response and the processing:
  - The gains need to be updated each time a change occur in the processing (LUT, algorithm, L1b calibration, etc...)
  - A strong effort of traceability in the gain computation, with respect to all other processing parameters, should be maintained.

## Vicarious Adjustment: Principle

- Computation starts from the decomposition:  $\rho_{gc}(\lambda) = \rho_{path}(\lambda) + t_d(\lambda) \cdot \rho_w(\lambda)$
- Knowing the true (or targeted) signal through in-situ measurements, individual gains are computed match-up per match-up by
  - $g_i(\lambda) = [\rho_{path}^t(\lambda) + t_d^t(\lambda) \cdot \rho_w^t(\lambda)] / [\rho_{path}(\lambda) + t_d(\lambda) \cdot \rho_w(\lambda)]$
- Average gains are finally derived:  $G(\lambda) = \text{mean}(g_i(\lambda))$
- To calibrate the NIR, one requires:
  - the most NIR band (865) is perfectly calibrated ( $\rightarrow$  *assumption*),
  - $\rho_w(\lambda_{NIR})$  is truly negligible, ( $\rightarrow$  *oligotrophic sites*)
  - aerosol spectral dependency is known ( $\rightarrow$  specific sites).
- Thus one has
  - $\rho_{gc}^t(865) = \rho_{gc}(865) = \rho_{path}(865)$
  - $\rho_{gc}^t(779) = g_i(779) \cdot \rho_{gc}(779) = \rho_{path}^t(779)$
- Using provides  $\rho_{path}^t(779)$  **hence  $g_i(779)$**
- Once NIR is calibrated,  $\rho_{path}^t(\lambda) = \rho_{path}(\lambda)$  and  $t_d^t(\lambda) = t_d(\lambda)$  at all bands, **hence  $g_i(\lambda)$**

# Vicarious Adjustment: implementation

- NIR calibration sites: **South Pacific Gyre and South Indian Ocean**, selected model MAR 90% (including continental background in the free troposphere and H<sub>2</sub>SO<sub>4</sub> in the stratosphere of fixed  $\tau_{550} = 0.025$  and 0.005 resp.)
- In-situ data: AAOT, BOUSSOLE and MOBY (foreseen : NOMAD)



Data selection:

- no flags within 25x25 px

Results in:

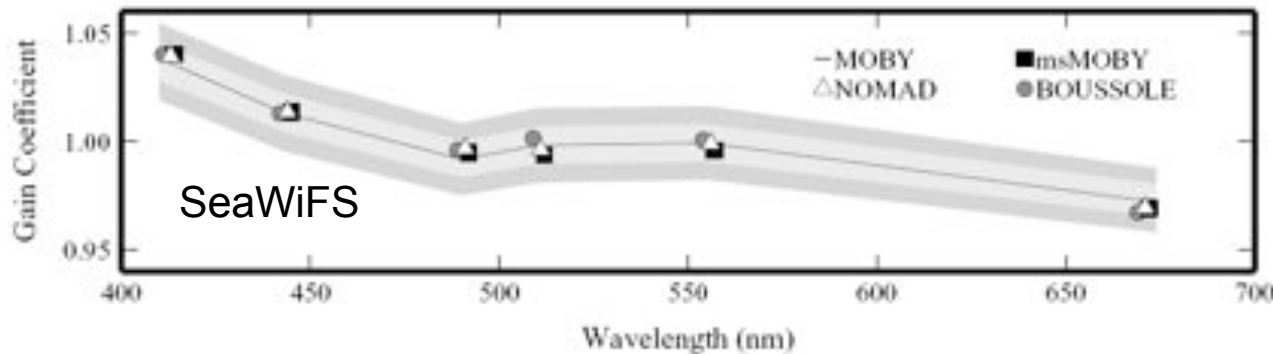
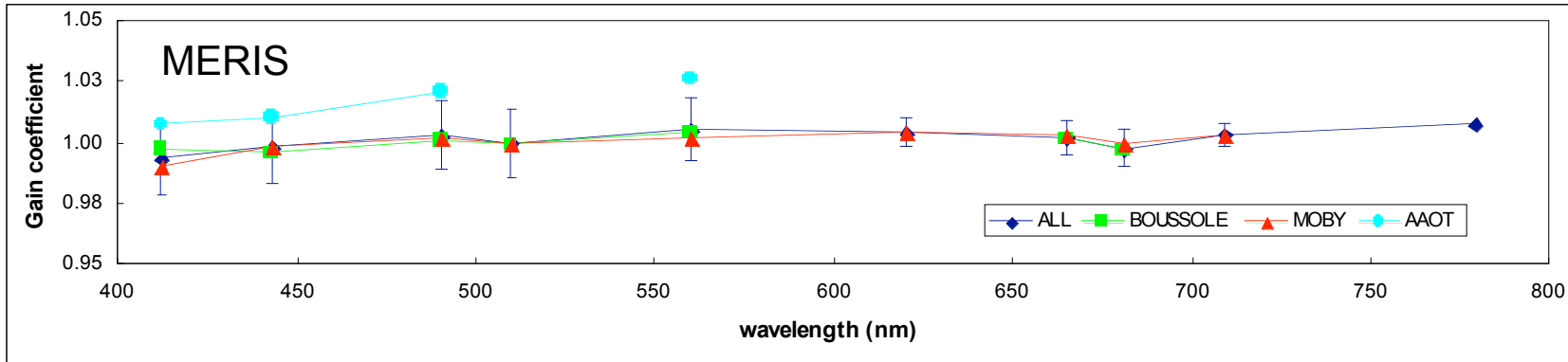
NIR (2003 & 2007 only):

- 14 at SPG
- 34 at SIO

VIS:

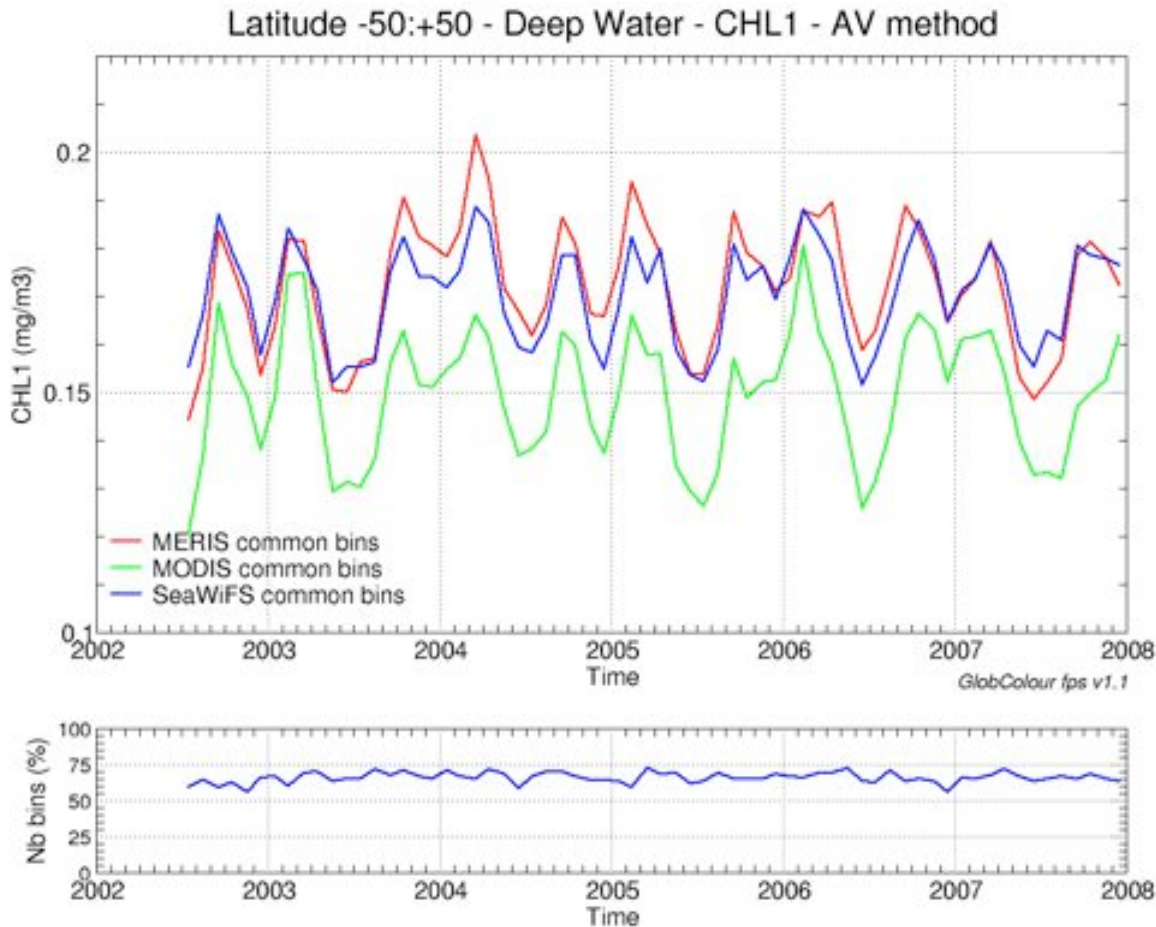
- 28 at MOBY
- 29 at BOUSSOLE
- 3 at AAOT

## Vicarious Adjustment: Preliminary Results



SeaWiFS figure from **Bailey S.W., S.B. Hooker, D. Antoine, B.A. Franz, P.J. Werdell**, « Sources and assumptions for the vicarious calibration of ocean color satellite observations », *Applied Optics*, 47(12), 2035-2045.

## INTER-SENSOR COMPARISONS: CHLa



**MERIS:**  
no vicarious  
adjustment

Source GlobColour,

Courtesy:

SeaWiFS: NASA (v5.2)

MODIS: NASA (v1.1)

MERIS: ESA (v2.0Q)