



*Revisiting Seasonal Plankton Cycles in the
Subarctic Atlantic and Pacific*

Ocean Color Research Team Meeting 2007

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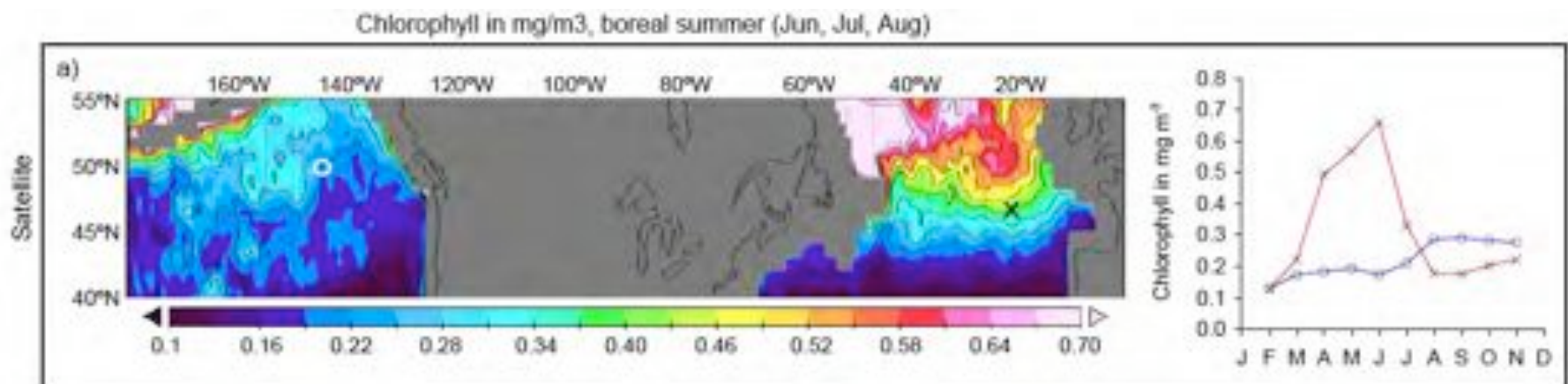
NOAA GFDL

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Seasonal Chlorophyll Cycles

- ✧ "In the North Atlantic, there is a marked increase in the standing stock of phytoplankton as measured by **Chlorophyll a** (...); in contrast, at Station P average values vary over a narrow range (...)" (Parsons & Lalli, 1988)
- ✧ "Based on the quantity of **Chlorophyll a** (...), there is virtually no seasonal variation in phytoplankton standing stock at Station P." (Frost, 1991)
- ✧ "**Chlorophyll a** at Station P is nearly constant year around (...) the key feature of the North Pacific is that the phytoplankton population is low and hardly changes with season." (Sarmiento & Gruber, 2006)



So what?

✧ Cellular pigmentation (Chl:C)
varies in response to

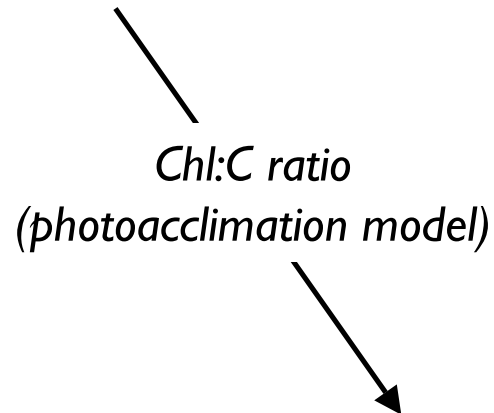
- ✧ Light
- ✧ Temperature
- ✧ Nutrient Limitation
- ✧ Iron Deficiency

✧ Subarctic Atlantic and Pacific
differ with respect to

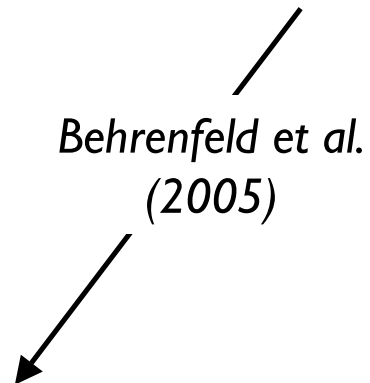
- ✧ Light
- ✧ Temperature
- ✧ Nutrient Limitation
- ✧ Iron Deficiency

Two Approaches

Satellite chlorophyll



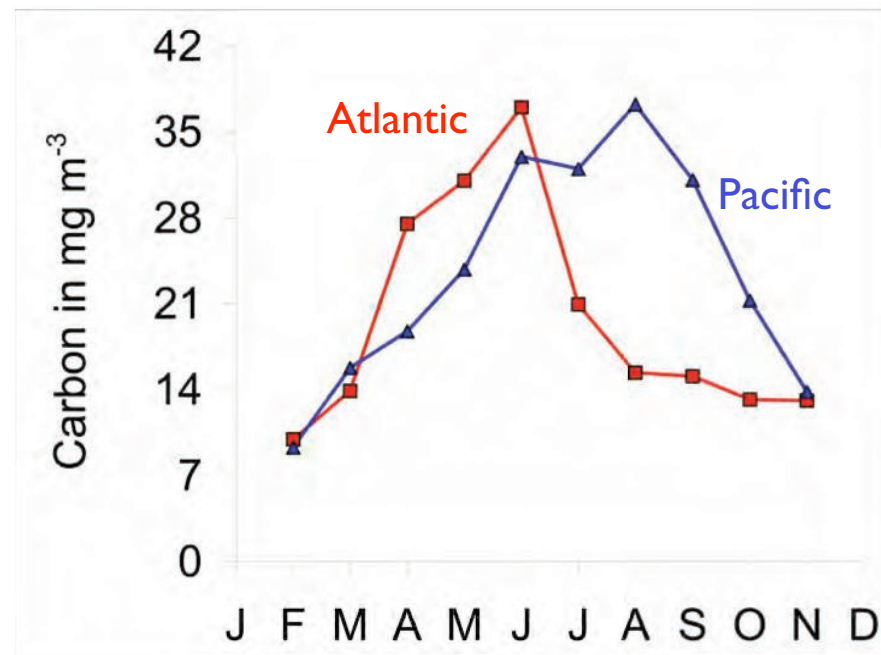
Satellite particulate backscattering



Phytoplankton carbon biomass

Conclusions

- ✧ Phytoplankton biomass in the Northeast Pacific varies seasonally.
- ✧ Maximum biomass concentrations are similar to the North Atlantic.



Phytoplankton Physiology

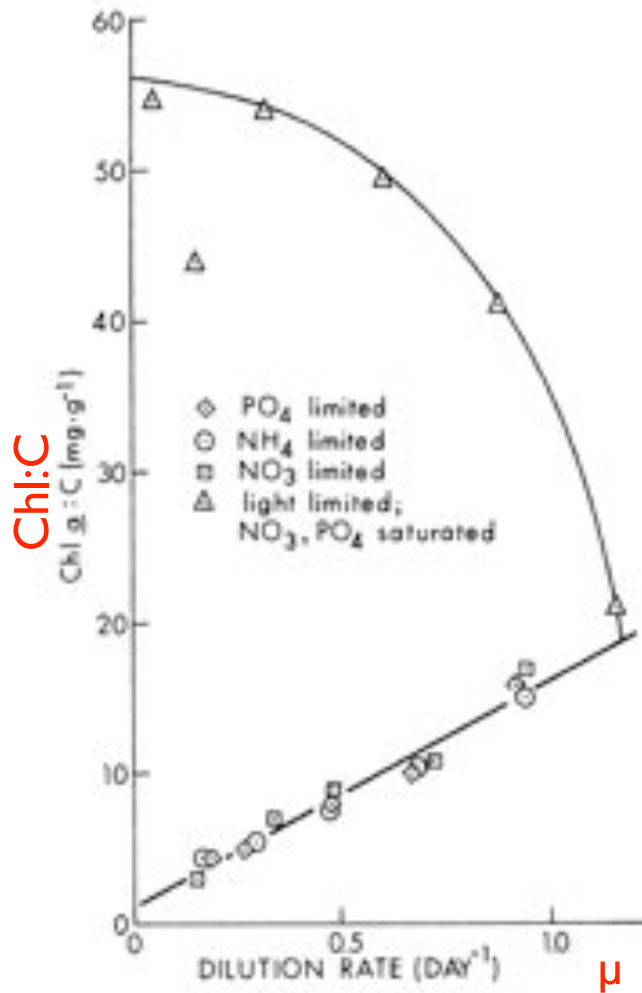
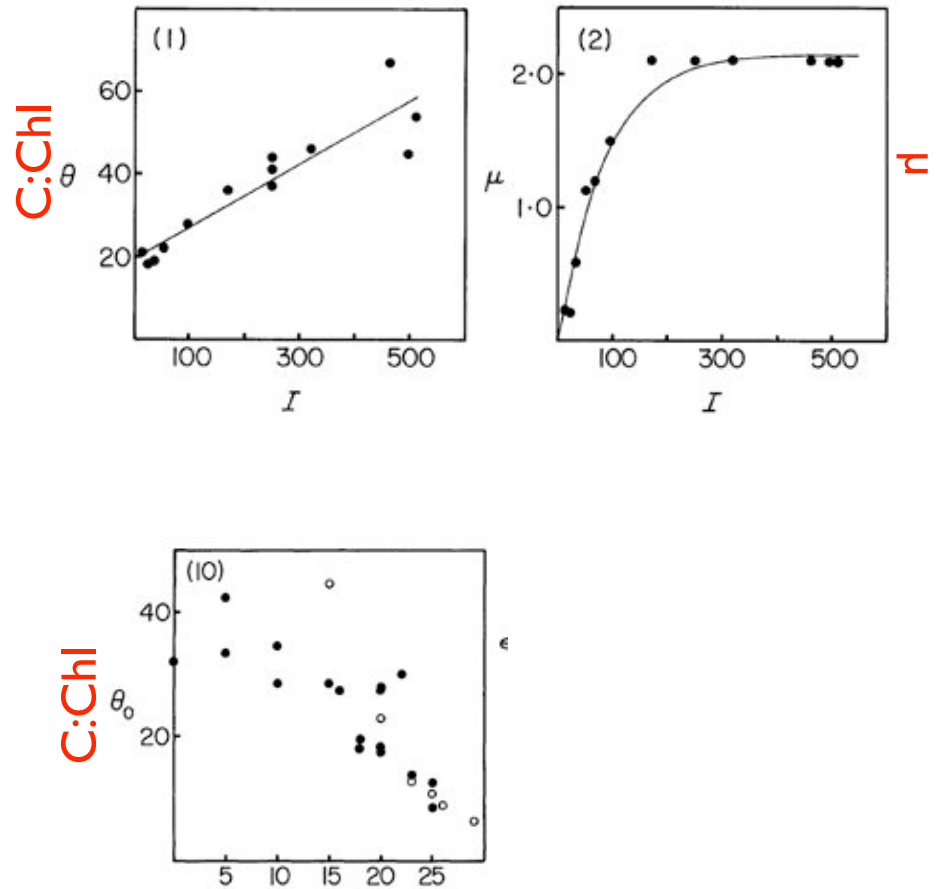


Fig. 5. Smooth curve to light-limited data from Bannister's model. Regression to nutrient-limited data, $\text{Chl } a:C = 1.14 + 15.1 \mu$.

Laws & Bannister (1980)
Light- and nutrient-limited growth of
Thalassiosira fluviatilis



R. Geider (1987)
Light- and temperature-limited growth
Thalassiosira pseudonana

Phytoplankton Physiology

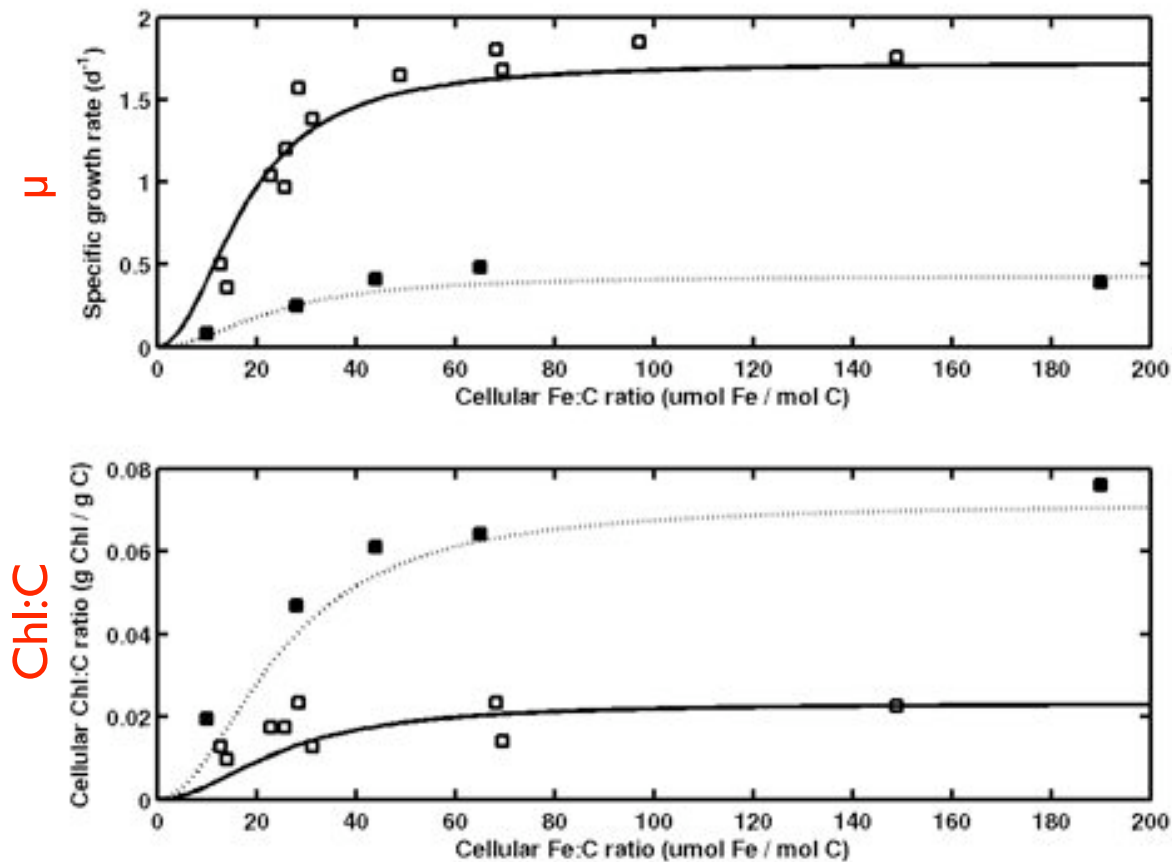


Figure S2 - Model fit to Sunda [1997]. Chl:C ratio (gChl gC^{-1}) and specific growth rate data (d^{-1}) for *Thalassiosira Pseudonana* under high light (open) and low light (filled) as a function of iron content. $P_m^C = 2.9\text{E-}05 \text{ s}^{-1}$, $\alpha = 3.5 \text{ gC gChl}^{-1} \text{ m}^2 \text{ Watts}^{-1} \text{ s}^{-1}$, $\text{Fe:N}_{\text{DFF}} = 25 \mu\text{molFe molC}^{-1}$, $T = 20 \text{ }^\circ\text{C}$, $\theta_{\text{max}} = 0.08 \text{ gChl gC}^{-1}$, $\kappa = 0.063 \text{ }^\circ\text{C}^{-1}$. Values were fit for θ_{max} , α , Fe:N_{DFF} and the low light irradiance ($24 \text{ Einstein m}^{-2} \text{ s}^{-1}$), the last of which was adjusted from the value given in Sunda [1997] ($50 \text{ Einstein m}^{-2} \text{ s}^{-1}$) for the best fit.

Sunda (1997)

Effect of iron on growth rate and Chl:C ratio of *Thalassiosira pseudonana*

Photoacclimation Model

- ✧ “Bastardized” version of GFDL global ocean biogeochemistry model
- ✧ Based on Geider (1997), but modified to account for
 - ✧ Co-limitation by nitrate, phosphate, silicate
 - ✧ The direct physiological effect of iron deficiency
 - ✧ Two size classes (“large” and “small”)

$$\vartheta_i = \vartheta_{\min,i} + \frac{(\vartheta_{\max,i} - \vartheta_{\min,i}) \cdot \frac{[Fe/N]_i^2}{k_{Fe} + [Fe/N]_i^2}}{1 + \vartheta_{\max,i}^{Fe} \cdot \alpha_i \cdot E / (2 \cdot P_{\max,i}^C \cdot \Lambda_{N,i} \cdot \exp(\kappa \cdot T))} \quad i \in [Lg, Sm]$$

Geider (1997), Dunne et al. (2006), Sunda (1997), Sunda & Huntsmann (1997)

Photoacclimation Model

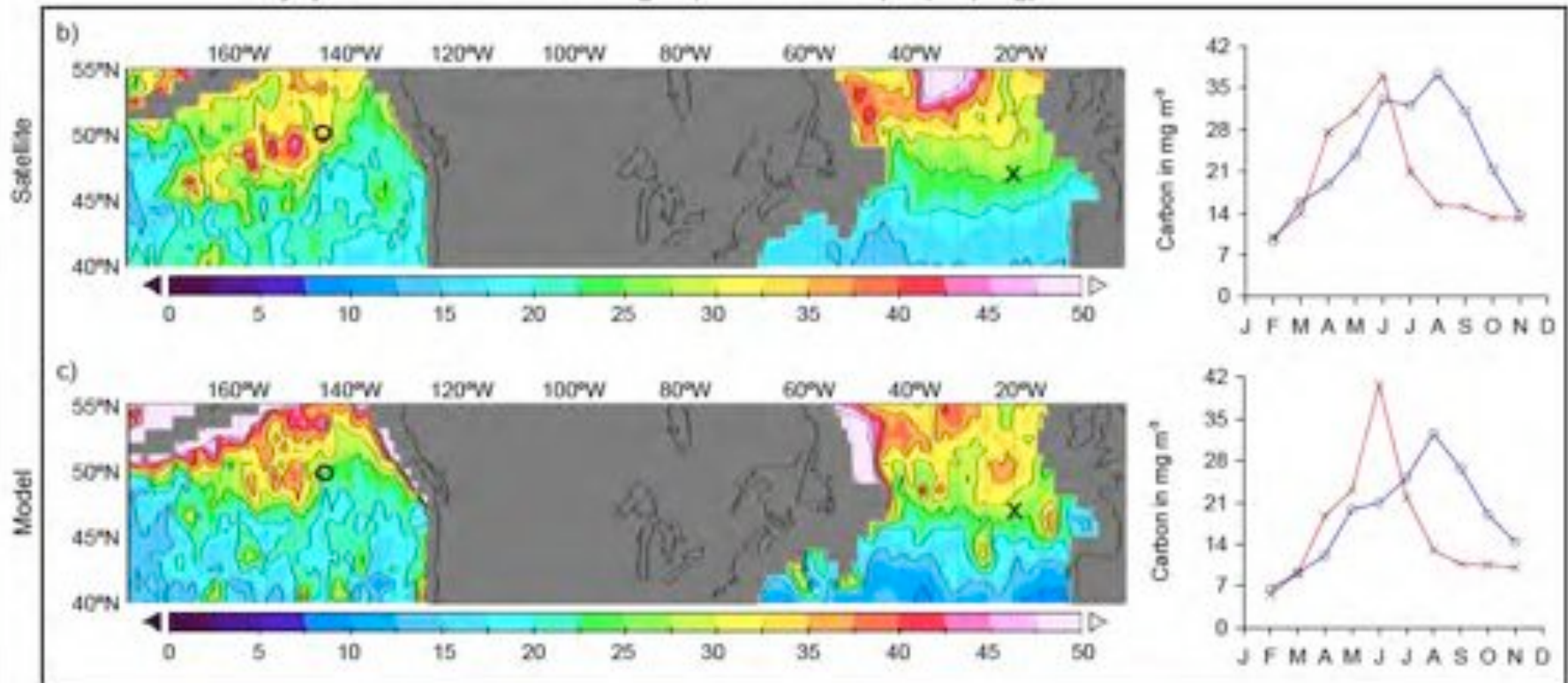
- ✧ “Bastardized” version of GFDL global ocean biogeochemistry model
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 - ✧ Co-limitation by nitrate, phosphate, silicate
 - ✧ The direct physiological effect of iron deficiency
 - ✧ Two size classes (“large” and “small”)
- ✧ Community Chl:C ratio = weighted average of “large” and “small” Chl:C

$$\vartheta = f_{Lg} \cdot \vartheta_{Lg} + (1 - f_{Lg}) \cdot \vartheta_{Sm}$$

Armstrong (1999), Dunne et al. (2005)

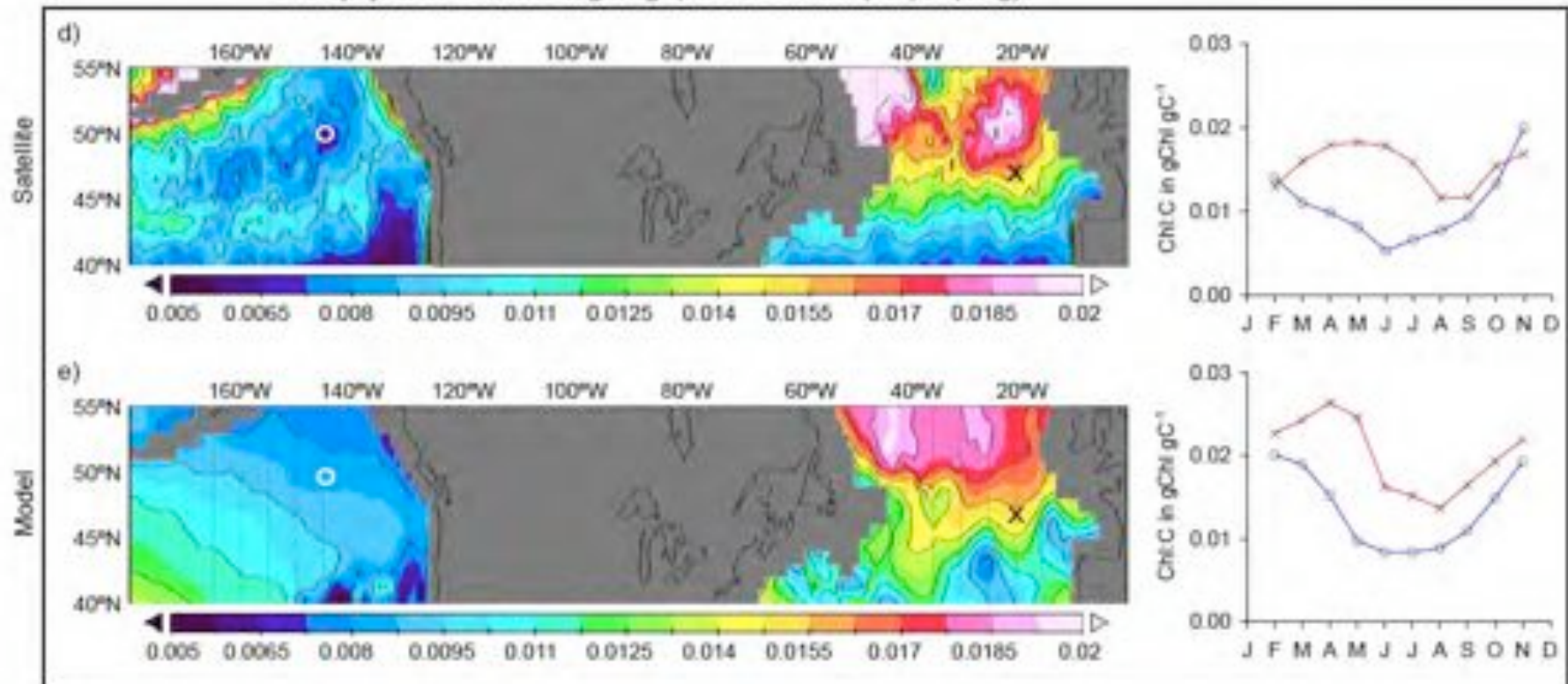
Phytoplankton Carbon Biomass

Phytoplankton carbon biomass in mg/m³, boreal summer (Jun, Jul, Aug)



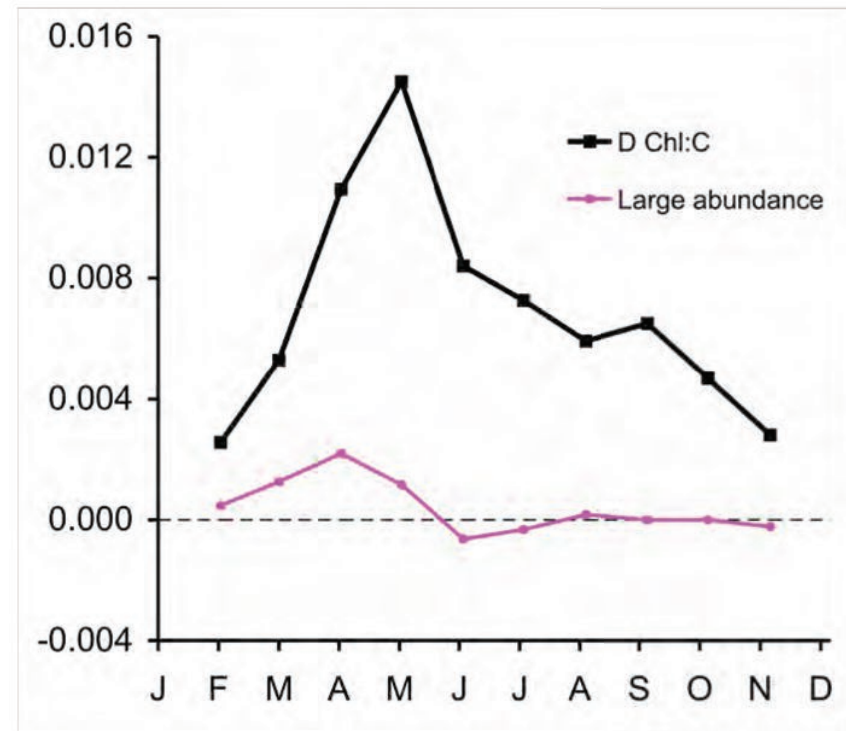
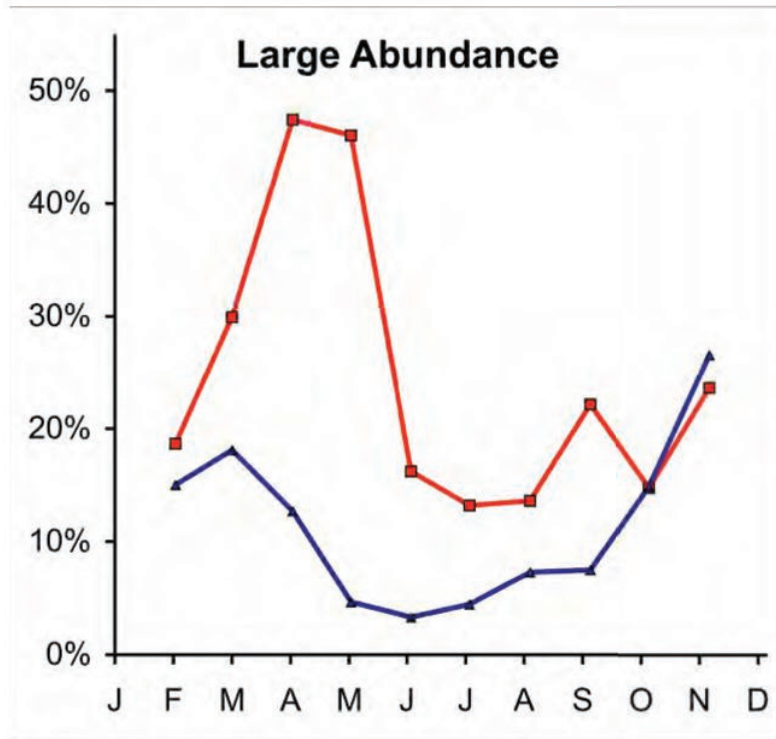
Chl:C Ratio

Chlorophyll to carbon ratio in gChl/gC, boreal summer (Jun, Jul, Aug)

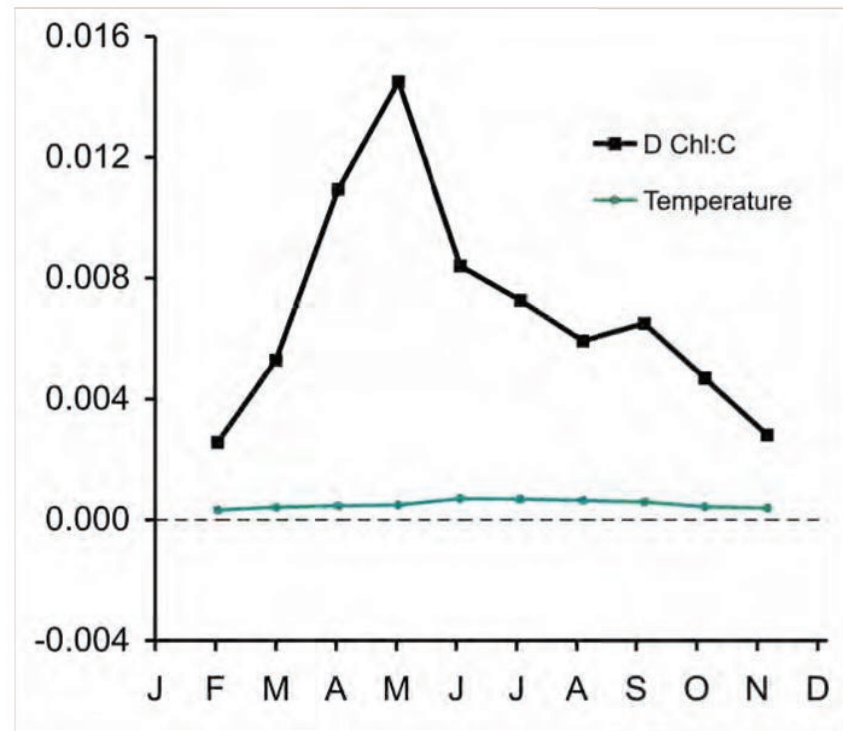
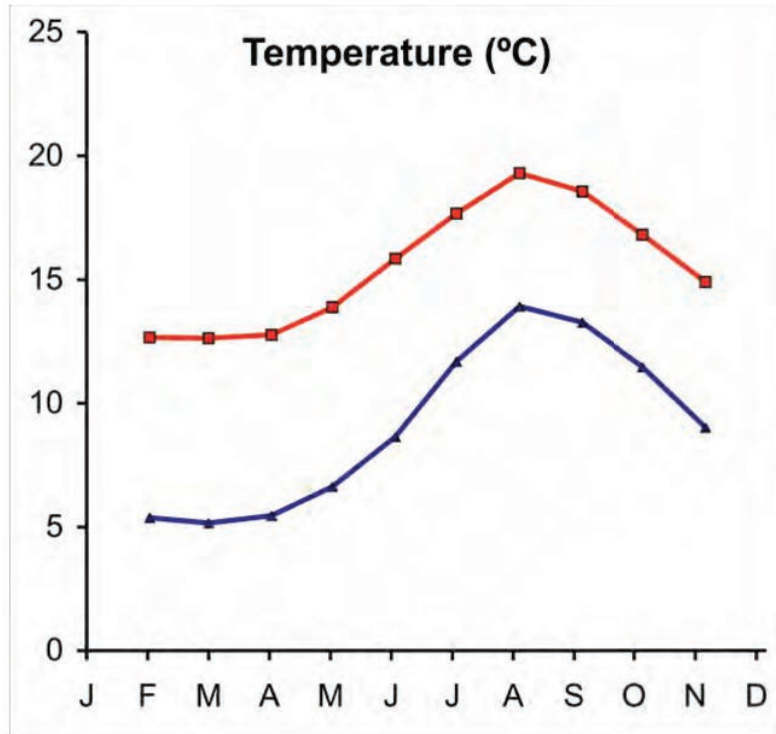


$$\Delta\vartheta \approx \sum_X \left(\frac{\partial\vartheta}{\partial X} \cdot \Delta X \right) \quad X \in [f_{Lg}, T, E, \Lambda_N, Q_{Fe}]$$

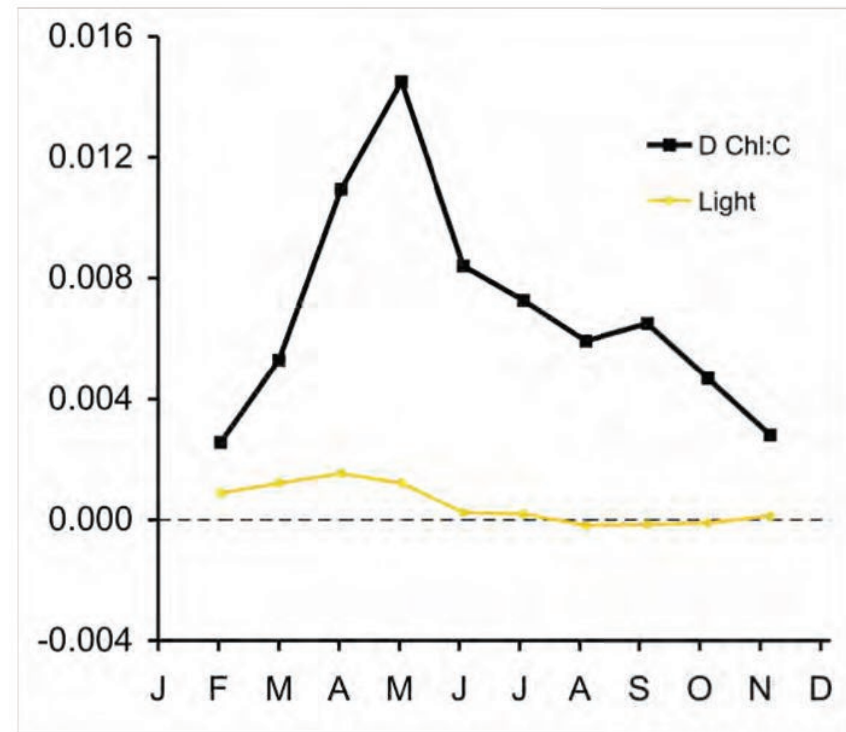
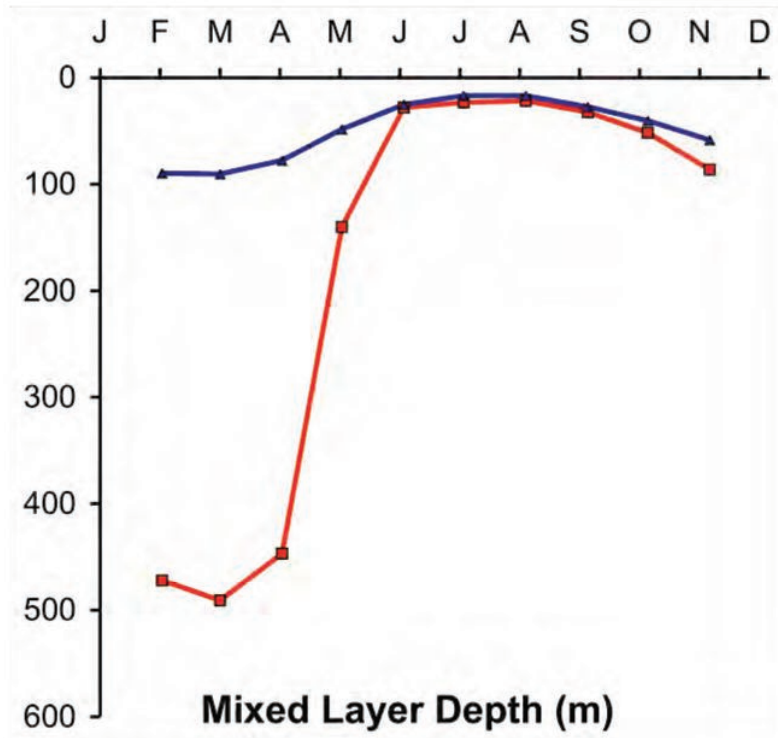
Decomposing the Chl:C Ratio



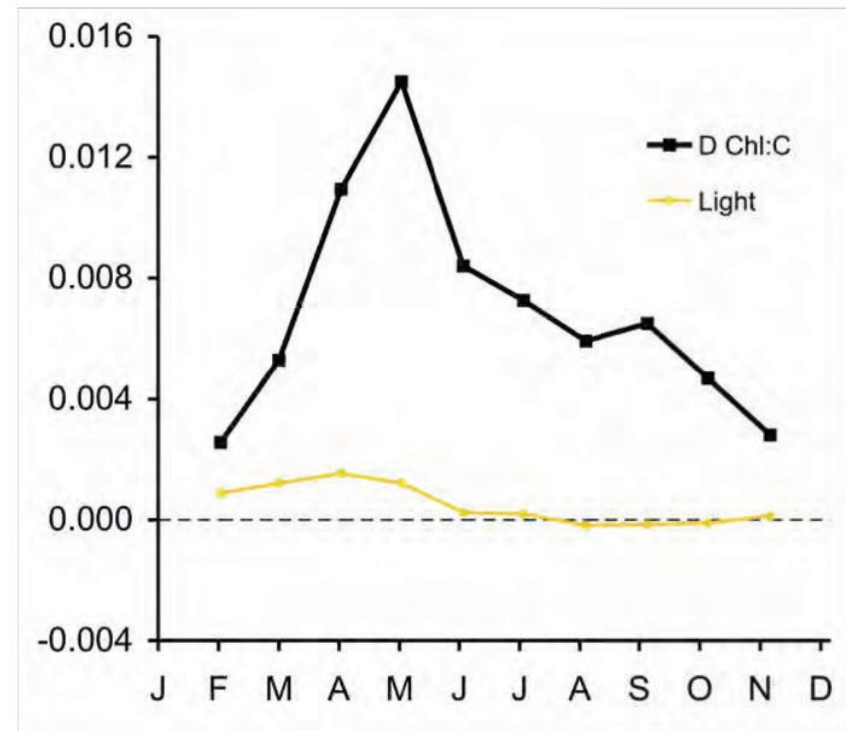
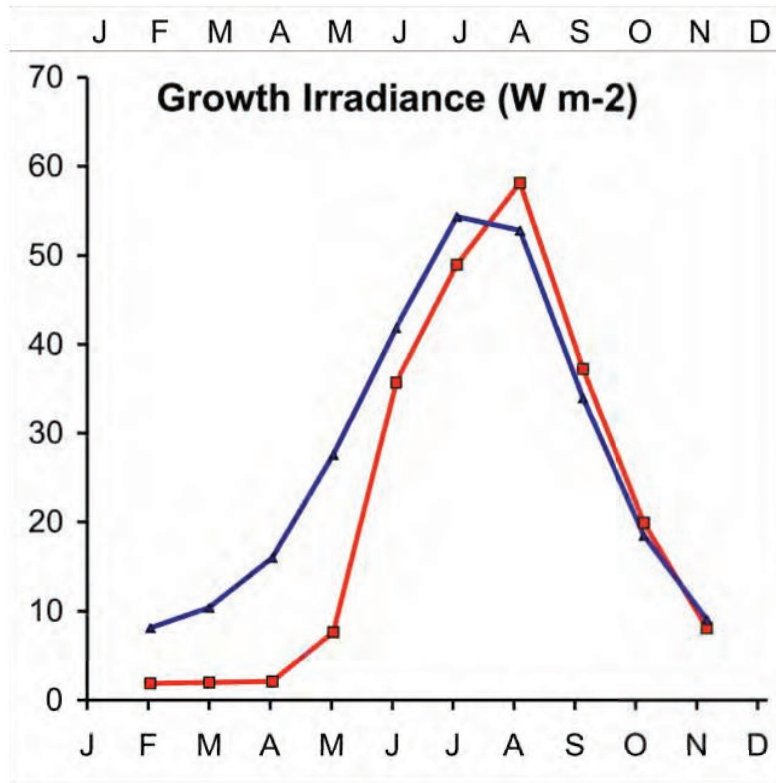
Decomposing the Chl:C Ratio



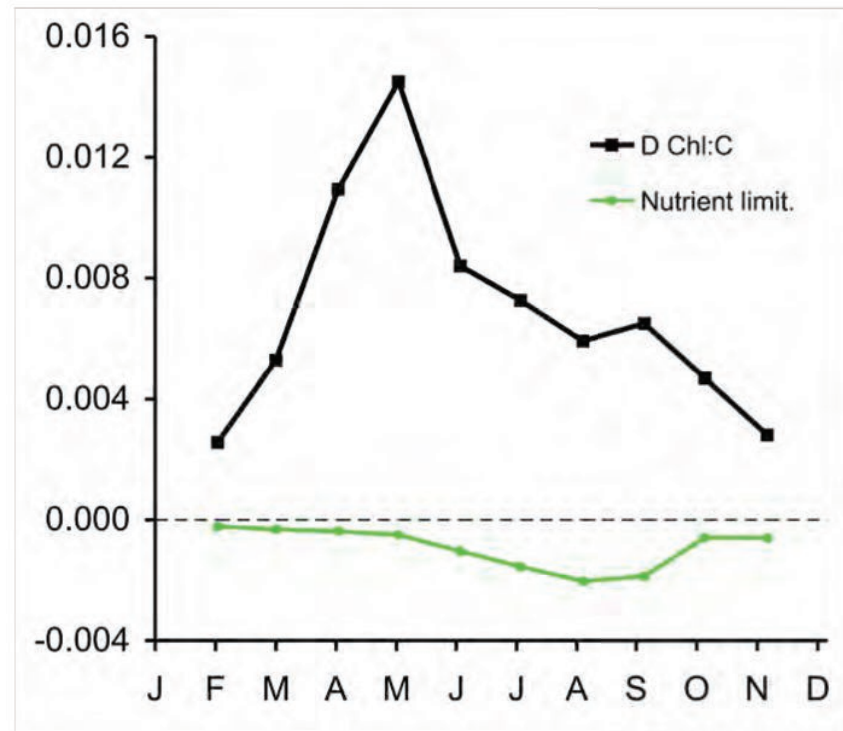
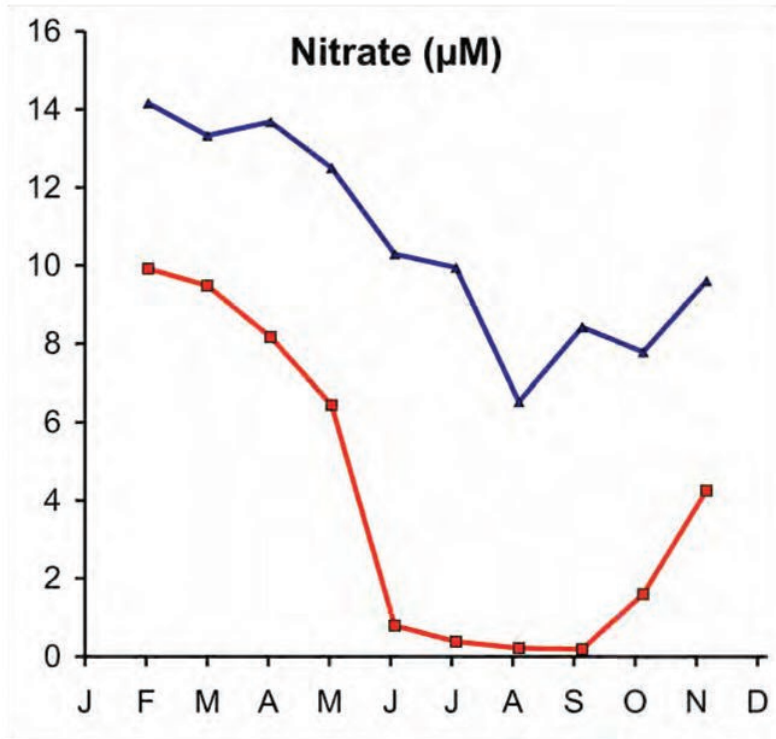
Decomposing the Chl:C Ratio



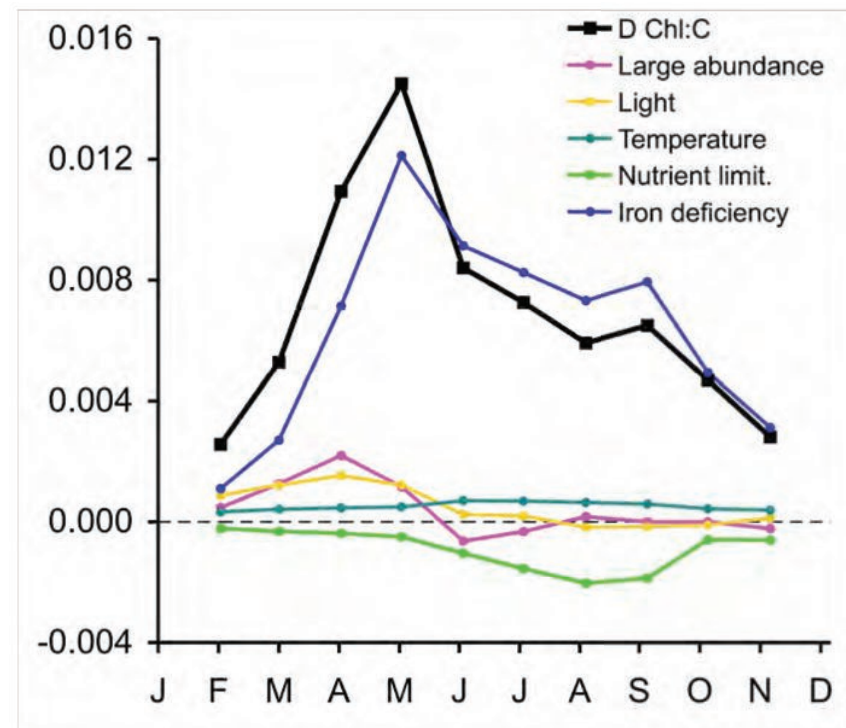
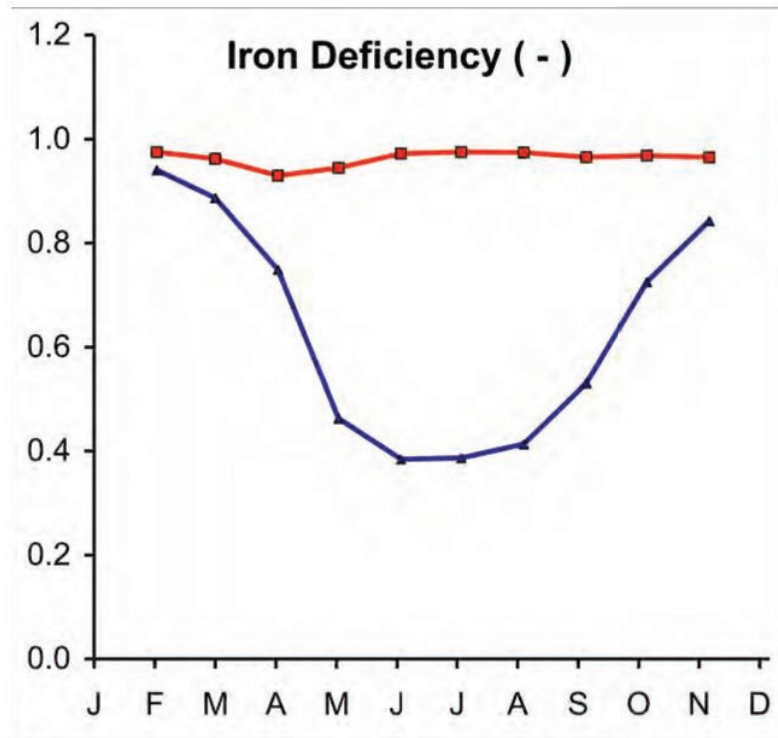
Decomposing the Chl:C Ratio



Decomposing the Chl:C Ratio



Decomposing the Chl:C Ratio



Phytoplankton Physiology

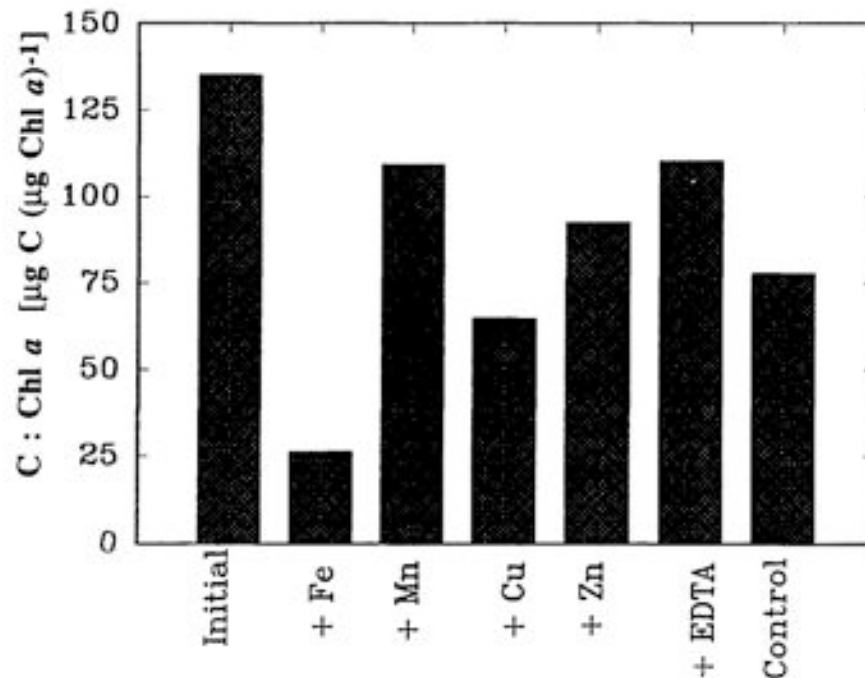
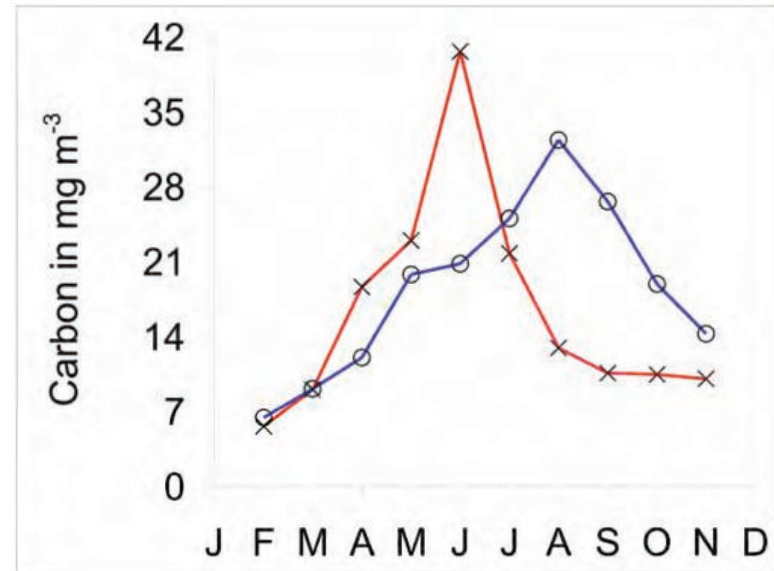


Fig. 6. Initial and final C: Chl *a* ratio for each metal treatment and control. Photosynthetic C was calculated to be the sum of the autofluorescent and diatom C as described in the text.

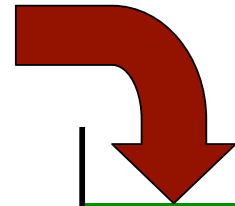
K. Coale (1991)

“Effects of Iron, Manganese, Copper, and Zinc Enrichments on Productivity and Biomass in the Sub-Arctic Pacific”

The Balance between Growth and Loss



Growth

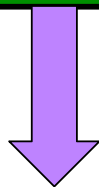


pe ratio
(Dunne et al.,
2005)

Sinking



Dilution



Non-sinking
mortality
(Grazing)

by difference

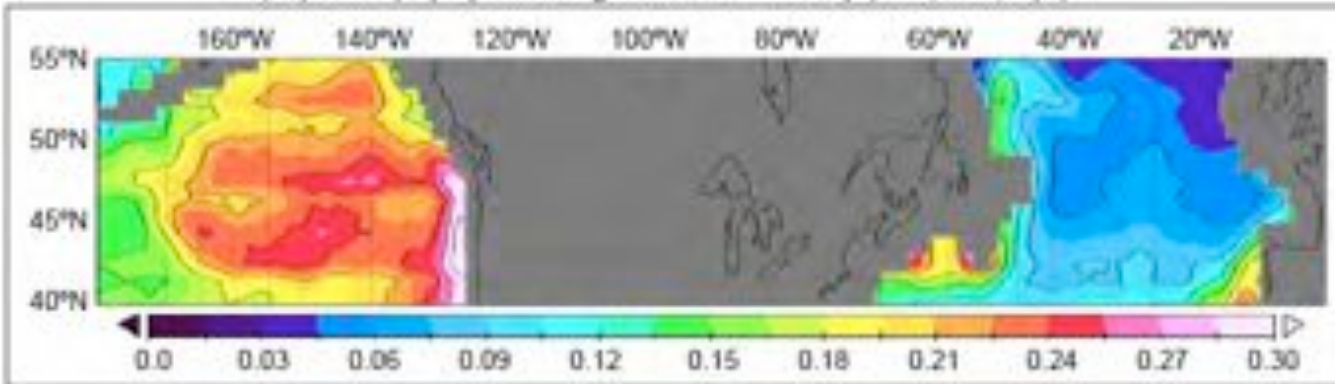


Standing Stock



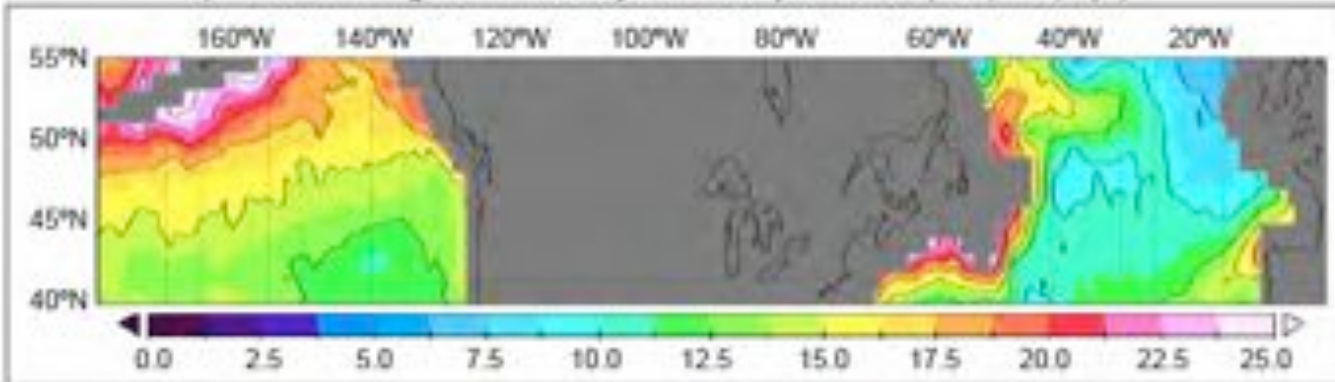
Net change

a) Specific phytoplankton growth rate in 1/day (Feb, Mar, Apr)



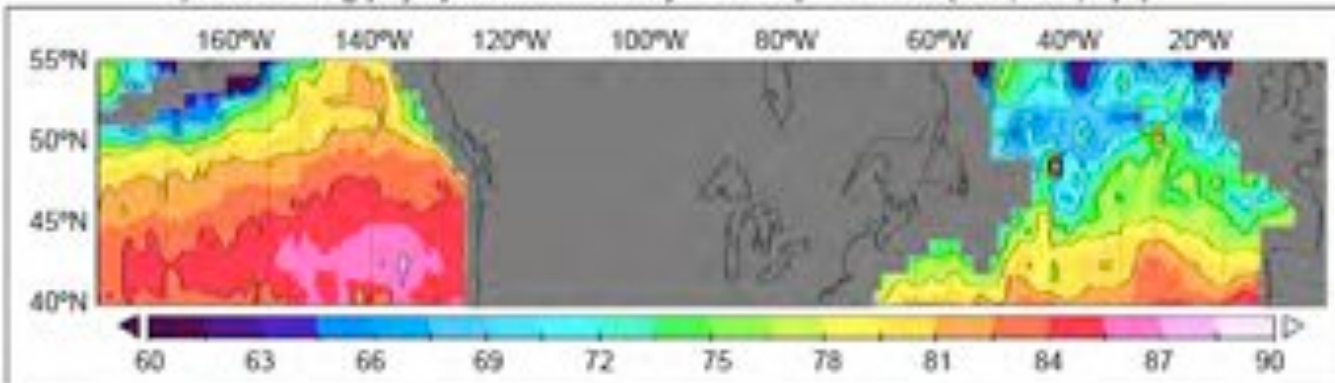
Growth rate

b) Particle sinking from mixed layer as % of production (Feb, Mar, Apr)



Particle export
(Dunne, 2005)

c) Non-sinking phytoplankton mortality as % of production (Feb, Mar, Apr)



Non-sinking
mortality
(grazing)

Conclusions (refined)

✧ North Atlantic

- ✧ Peak biomass and max. Chl:C coincide, leading to the quintessential chlorophyll bloom
- ✧ Despite low growth rates, biomass accumulates rapidly in spring as the result of low sinking and grazing losses.

✧ North Pacific

- ✧ Biomass increases are accompanied by decreases in Chl:C.
- ✧ Low chlorophyll concentrations reflect suppressed Chl:C, not low biomass.
- ✧ Iron deficiency is the main factor for reducing Chl:C.
- ✧ Despite higher growth rates, biomass accumulates less rapidly in the Pacific as the result of higher sinking losses and grazing pressure.