

Measuring and scaling net community production and carbon export in the Southern Ocean

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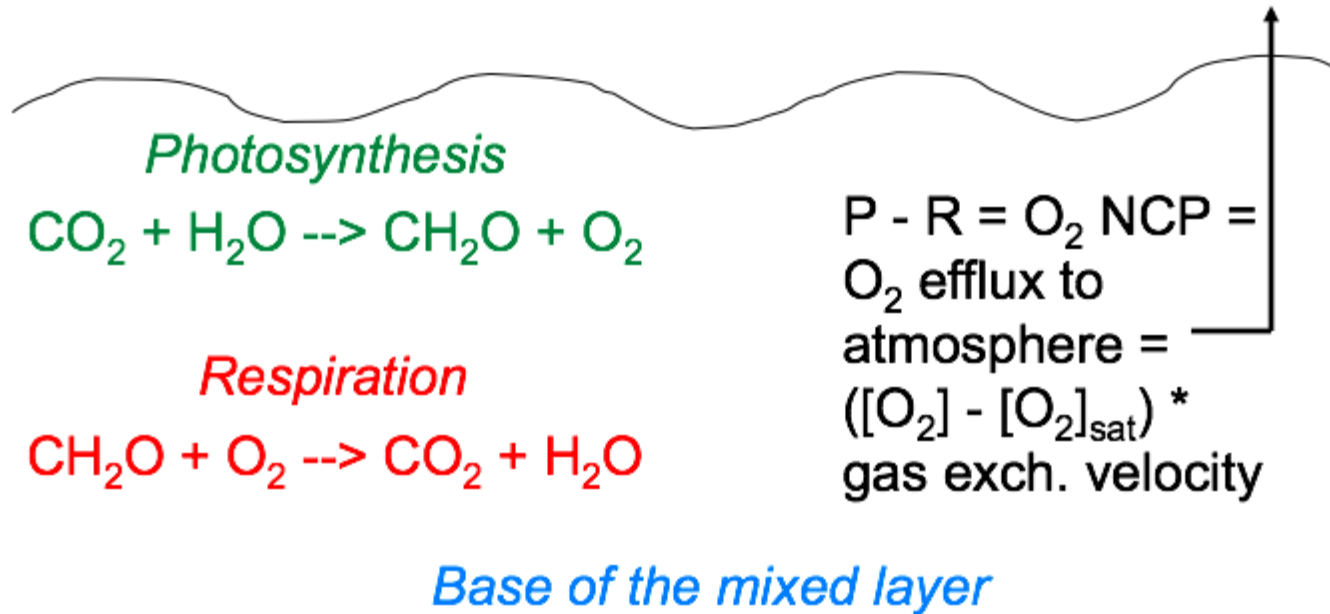
With experimental contributions and ideas from Nicolas Cassar

Many collaborators (especially Pedro Monteiro and Warren Joubert,
and Bronte Tilbrook)

Outline

- Basis of net community production measurements by O_2/Ar
- Relation between chlorophyll-normalized NCP ($mmol O_2 mg chl^{-1} day^{-1}$) and average irradiance in the mixed layer
- Strategy for scaling up NCP measurements
- Other methods for accessing NCP and carbon export
 - Summary of methods
 - Global NCP variations estimated from different approaches

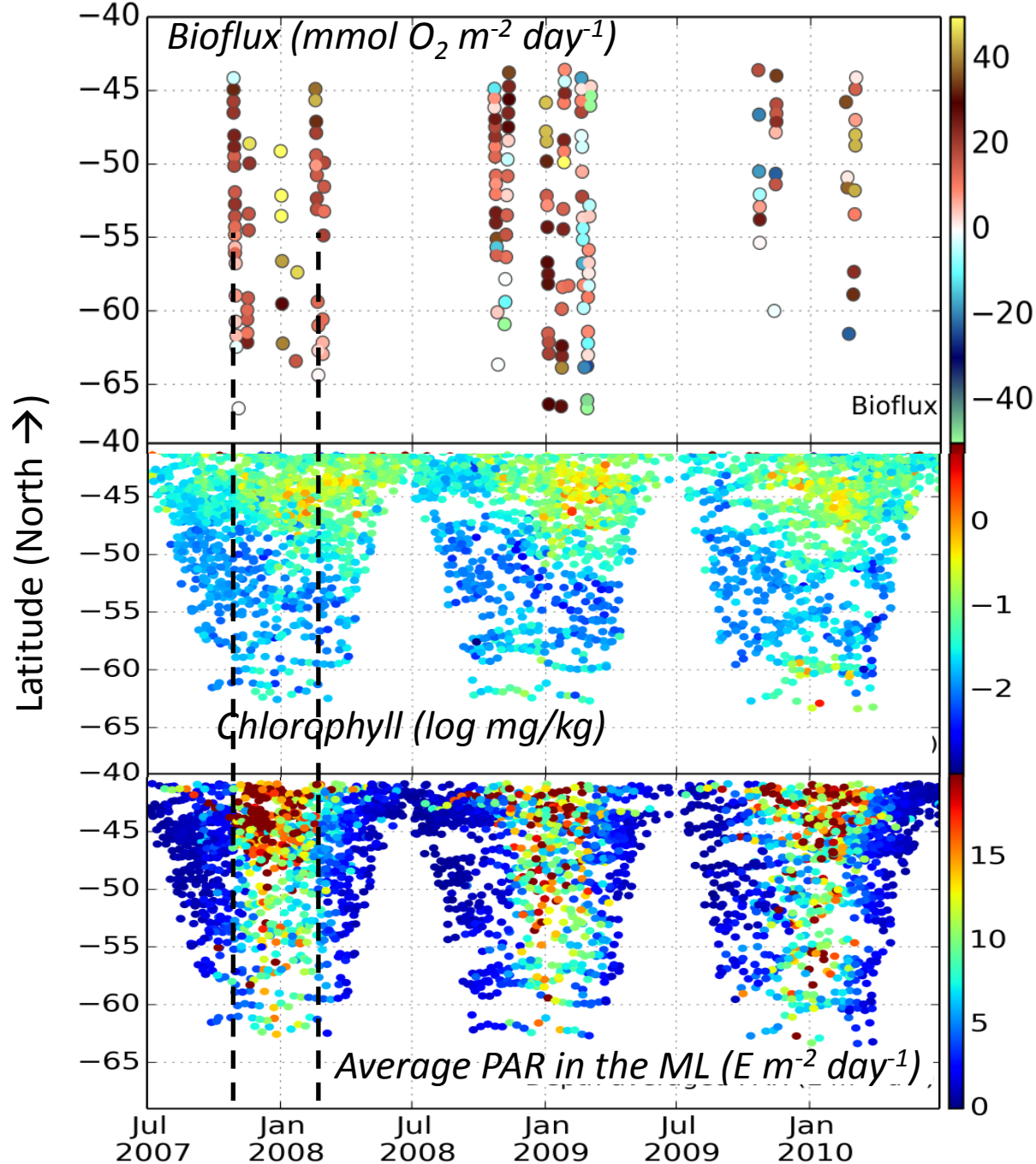
Mixed layer NCP and carbon export estimated from mixed layer O₂/Ar ratios



- In summer:
 - Photosynthesis > respiration
 - $[\text{O}_2] > [\text{O}_2]_{\text{sat}}$, lost to atmosphere
 - $\text{O}_2 \text{ efflux} = ([\text{O}_2] - [\text{O}_2]_{\text{sat}}) * \text{gas exchange velocity}$
- Net community O₂ production = flux to atmosphere
- Complication: $[\text{O}_2] > [\text{O}_2]_{\text{sat}}$ because of warming and bubble entrainment
- Measure Ar as inert analog to O₂ to correct for physical supersaturation (Jenkins, Quay, Emerson, Luz...)

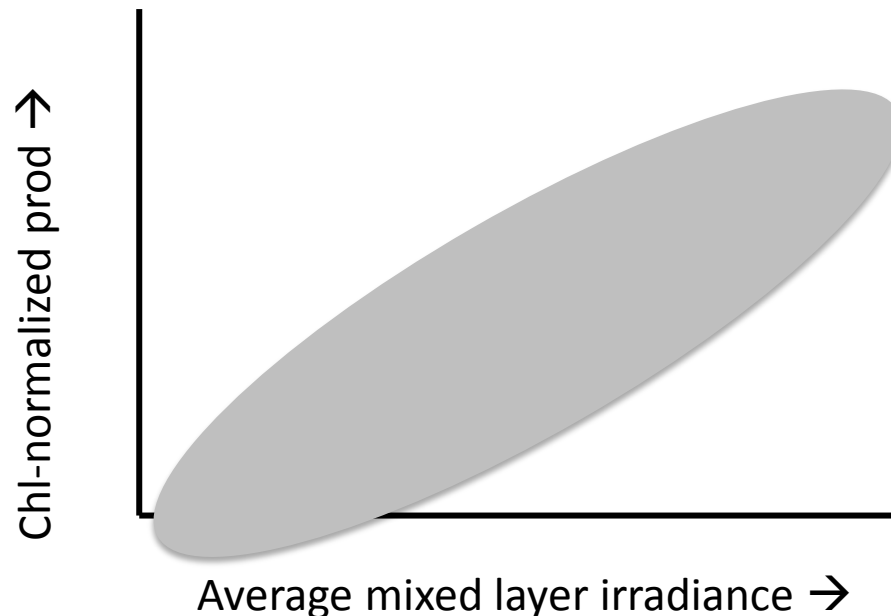
Net community production south of Australia

- Work on Astrolabe (French supply ship transiting between Hobart and Dumont d'Urville)
- Main collaborator: Bronte Tilbrook
- Data sources:
 - Bioflux from our O_2/Ar measurements
 - ML depth from Argo floats
 - Chl from satellite



Searching for a compact and scalable relation between O_2 bioflux (\sim NCP), MLD, chlorophyll, sea surface PAR and extinction coefficient

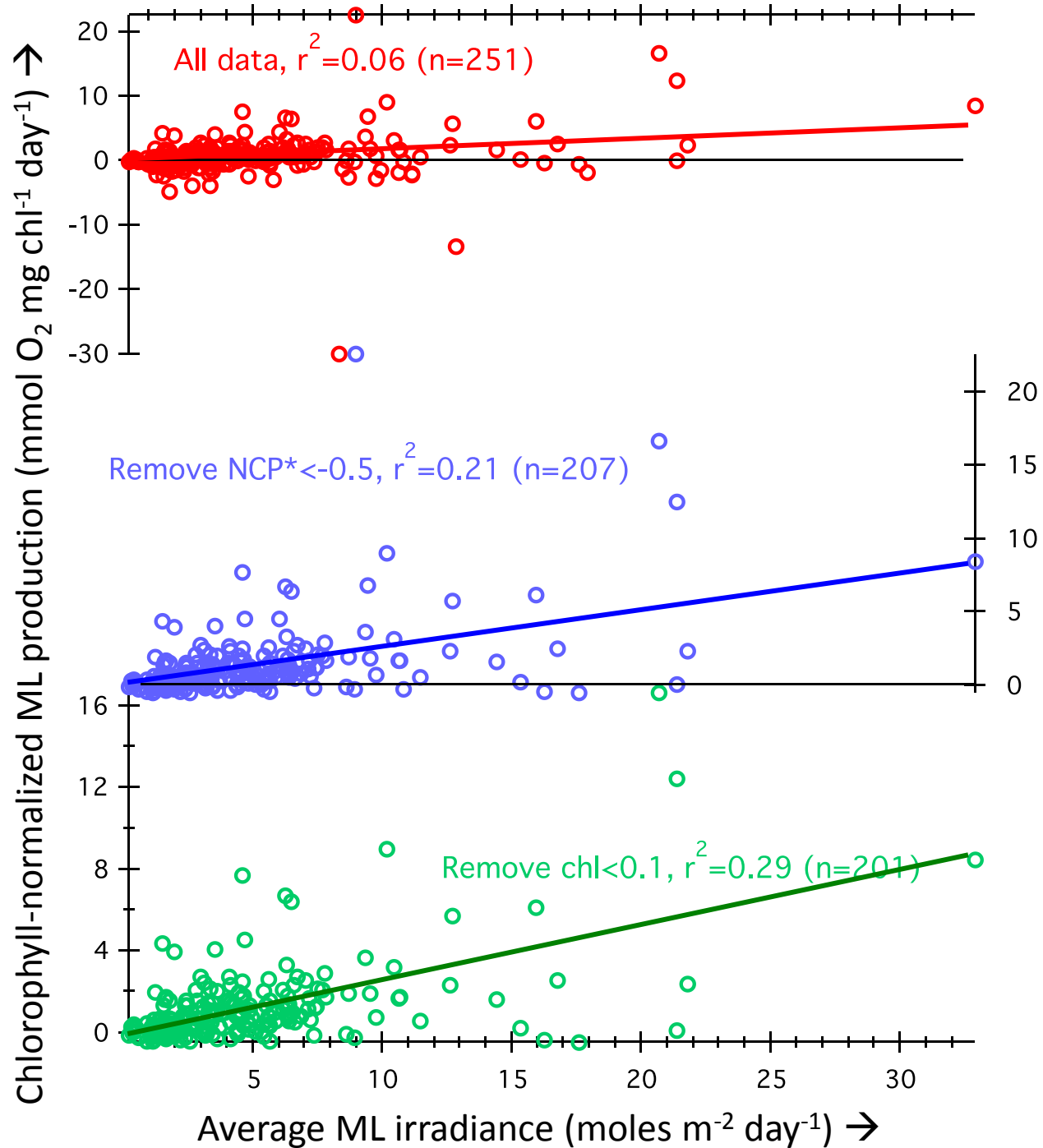
- Average mixed layer irradiance = $\int \text{Sea Surface PAR} * \exp(-kz) dz / \text{MLD}$
 - Units: $E m^{-2} day^{-1}$
 - K = extinction coefficient
- Chlorophyll-normalized production in ML = O_2 bioflux / (MLD*chl)
 - Units: $mmol O_2 mg chl^{-1} day^{-1}$



Data sources for plots of chlorophyll-normalized volumetric production vs. average ML irradiance

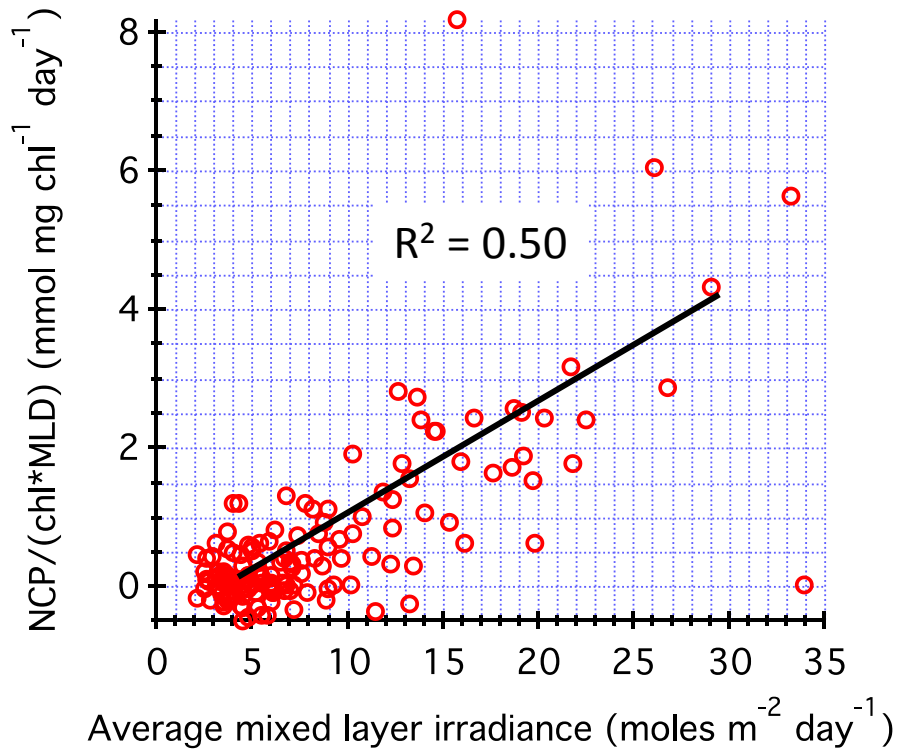
- O_2/Ar , chlorophyll, and mixed layer depth are measured
 - Satellite chlorophyll for Astrolabe
- Winds, sea surface PAR, and extinction coefficient are remote-sensing products
 - Morel equation for Astrolabe

Astrolabe data
(Hobart to Antarctica):
integrate and scale

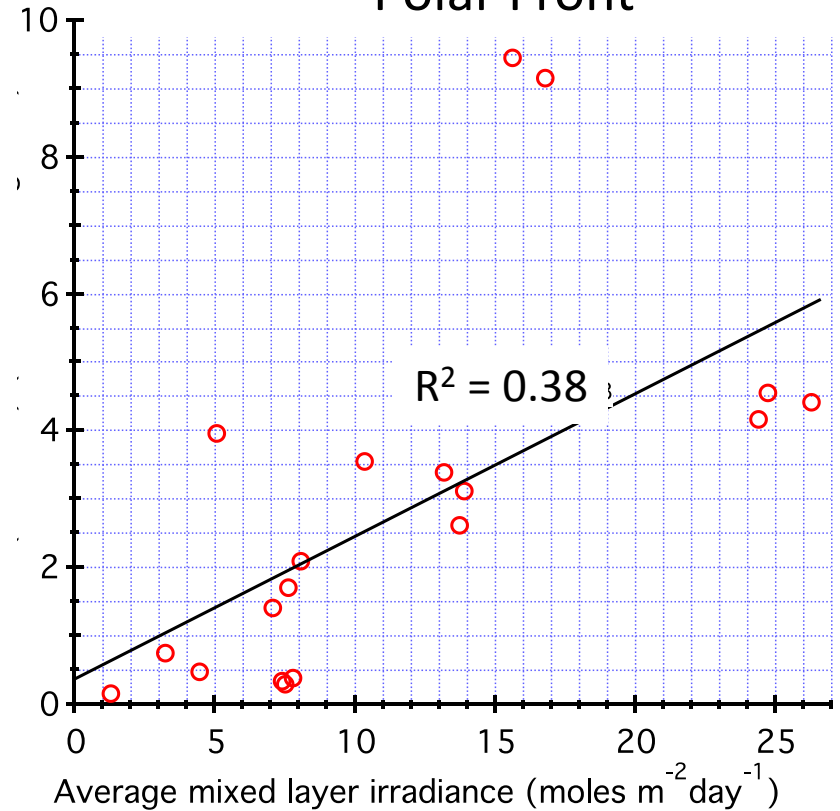


Agulhas and SAZ-SENSE data

Agulhas: Cape Town to Antarctica



SAZ-SENSE: Hobart south below Polar Front

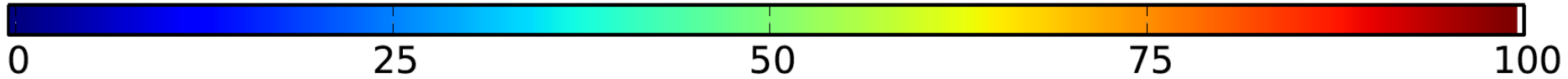


Data sources for simulated NCP field

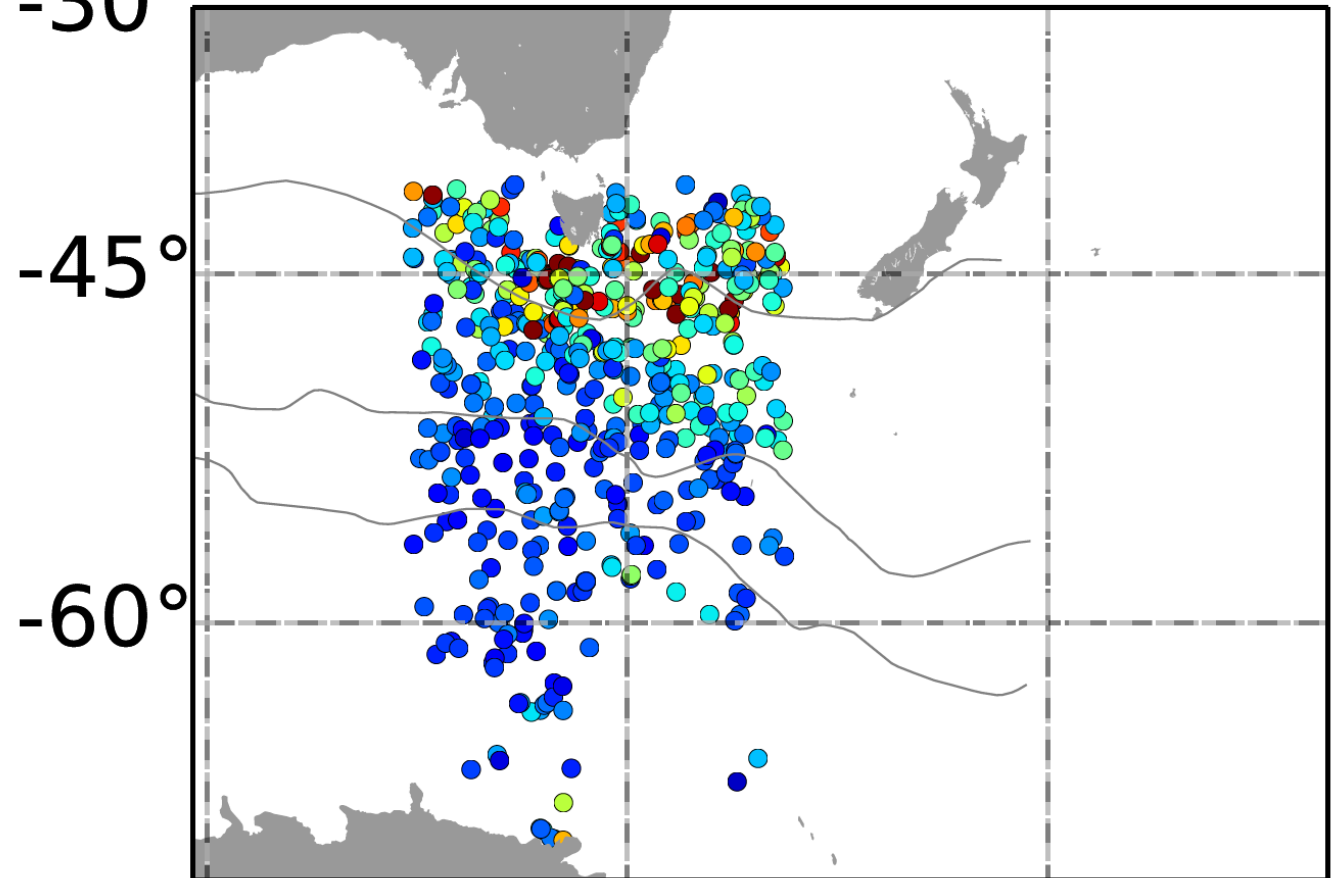
- Chl, sea surface PAR, extinction coefficient from satellite product
- MLD from “colocated” Argo float
- NCP calculated from the relationship between chl-normalized volumetric NCP and averaged ML irradiance

Scaling NCP in the Australian sector

Color indicates O_2 NCP ($\text{mmol m}^{-2} \text{day}^{-1}$)



-30° Feb, $k=0.36$, $m=0.19$, $r^2 = 0.45$

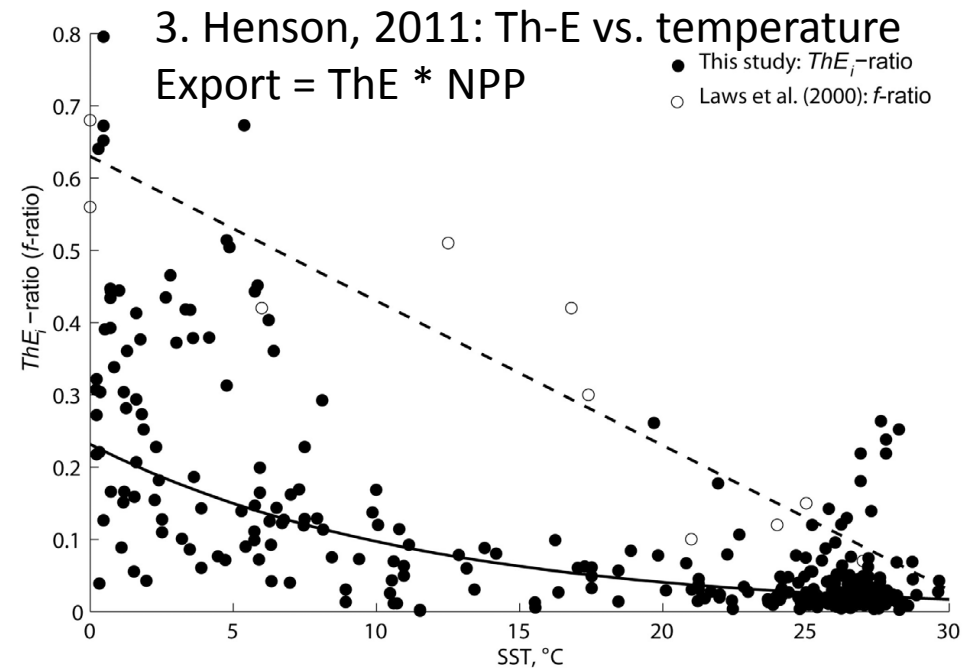
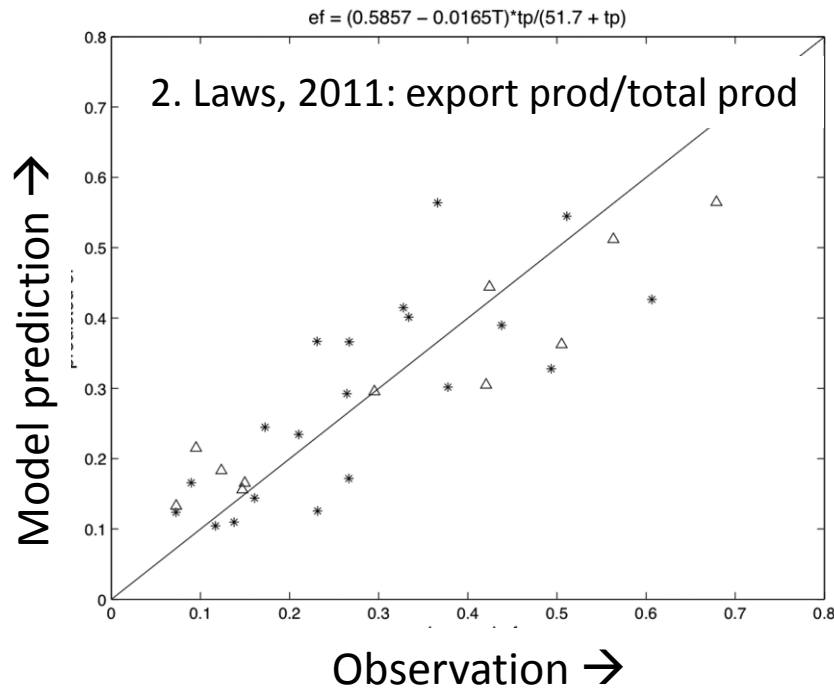
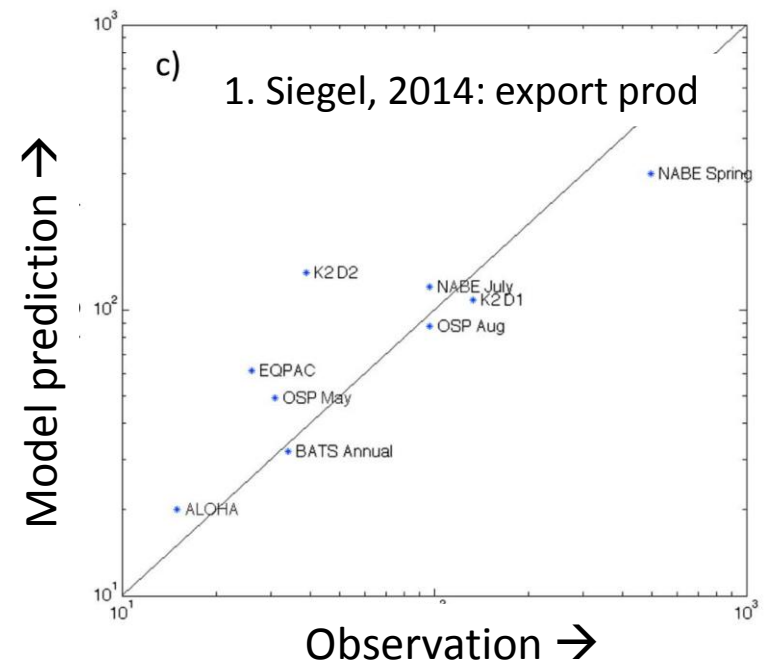


Methods for measuring net community production and carbon export

- Day to month scale
 - O_2/Ar : Uncertainties due to vertical mixing, gas exchange estimate; ML only
 - $^{15}NO_3^-$ assimilation: bottle effects, depends on day's irradiance; ML or EZ
 - $^{234}Th / C$ ratios: ML or EZ or arbitrary control depth
- Seasonal time scale
 - DIC or NO_3^- drawdown: lateral and vertical mixing; ML

4 recent statements about the global-scale pattern of variability NCP and export

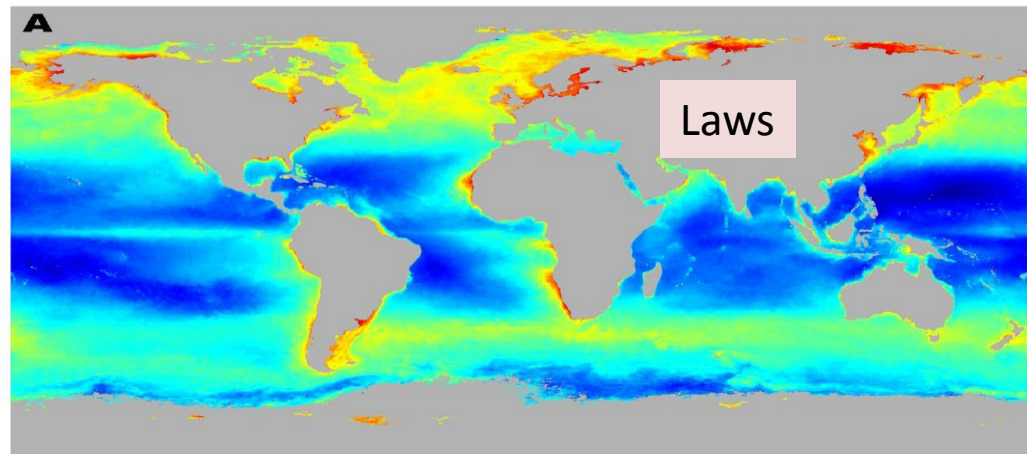
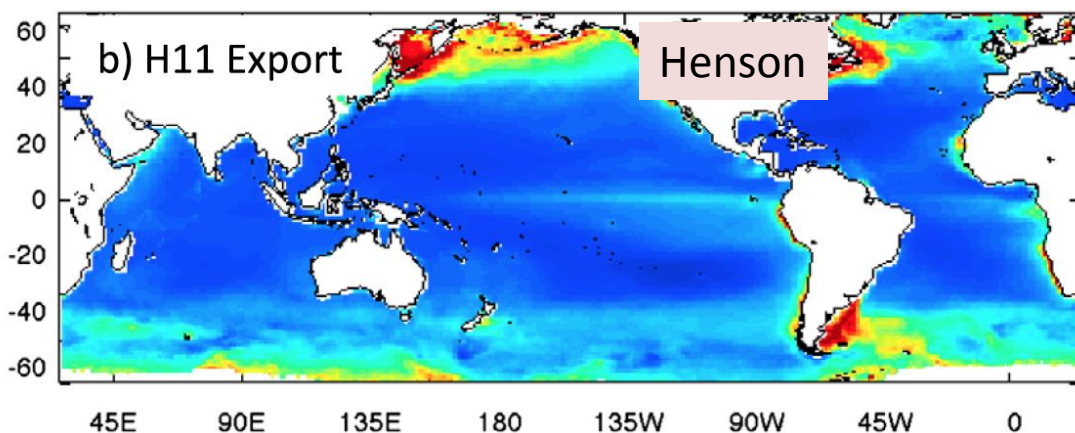
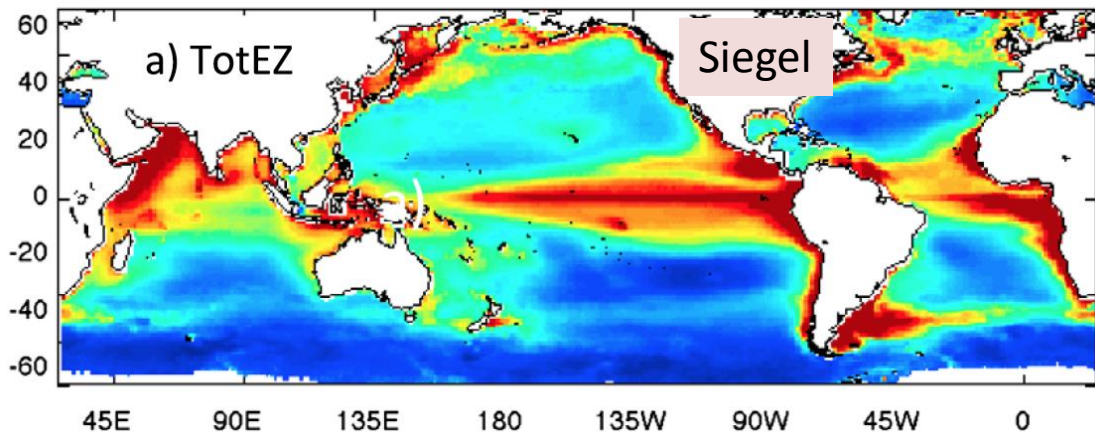
- Laws et al: simple model based on temperature dependence of photosynthesis and respiration
- Siegel et al: Simple 1-D ecosystem model
- Henson et al: empirical model relating The e-ratio to temperature



4. New production estimates from O₂/Ar or DIC balance show little variability

- Annual NCP based on ML O₂ or DIC balance (Emerson, 2014): NCP is in the range of 1-4 moles m⁻² yr⁻¹ in the
 - Southern Ocean
 - Equatorial Pacific Ocean
 - Subtropical north and south Atlantic
 - Subtropical North Pacific and western South Pacific
 - Subpolar North Pacific and North Atlantic

Inferred global patterns of new production or carbon export



Comments and conclusions

- Plots of chl-normalized NCP vs. average ML irradiance look like good candidates for scaling NCP observations
- It matters what depth domain NCP and carbon export are reported (e. g., ML or euphotic zone)
- Models should report ML carbon fluxes in addition to fluxes for other domains
- Extensive remineralization in the lower euphotic zone will lead to increasing DIC, decreasing O_2

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- Global models of upper ocean carbon fluxes should be tested vs. atmospheric O_2/N_2 data

Atmospheric O_2 concentration →

